



# ***Detection of the Earth's time-variable gravity field using GRACE***

***R. Schmidt, F. Flechtner, J. Kusche***

***GeoForschungsZentrum Potsdam (GFZ)***



# Outline

- The GRACE mission
- Overview results for static and time-variable GRACE gravity field
- Evolution of GRACE gravity product quality
- Currently open issues
- Conclusions



# The GRACE Mission

## GRACE = Gravity And Climate Recovery Experiment

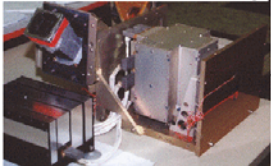
BlackJack GPS receiv.



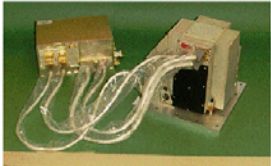
K-Band Horn



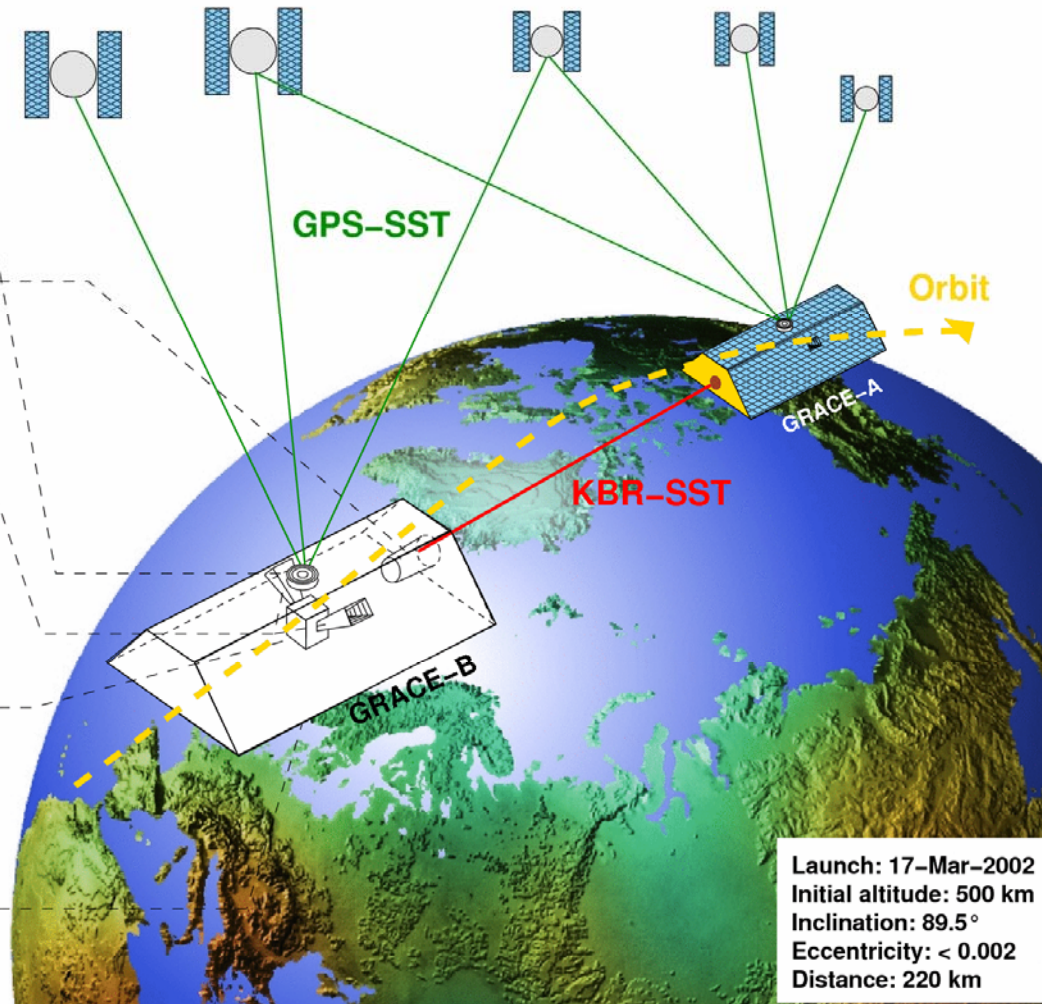
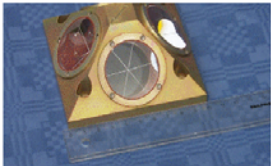
Star Camera Assembly



SUPERSTAR Acceler.



Laser Retro Reflector



- **NASA/DLR-Mission:**

Selected in 1997 within NASA's *Earth System Science Pathfinder (ESSP)* project

- **Principle Investigator (PI) and Co-PI:**

Byron Tapley (CSR),  
Christoph Reigber (GFZ,  
now Markus Rothacher)

- **US-German Science Data System (SDS):**

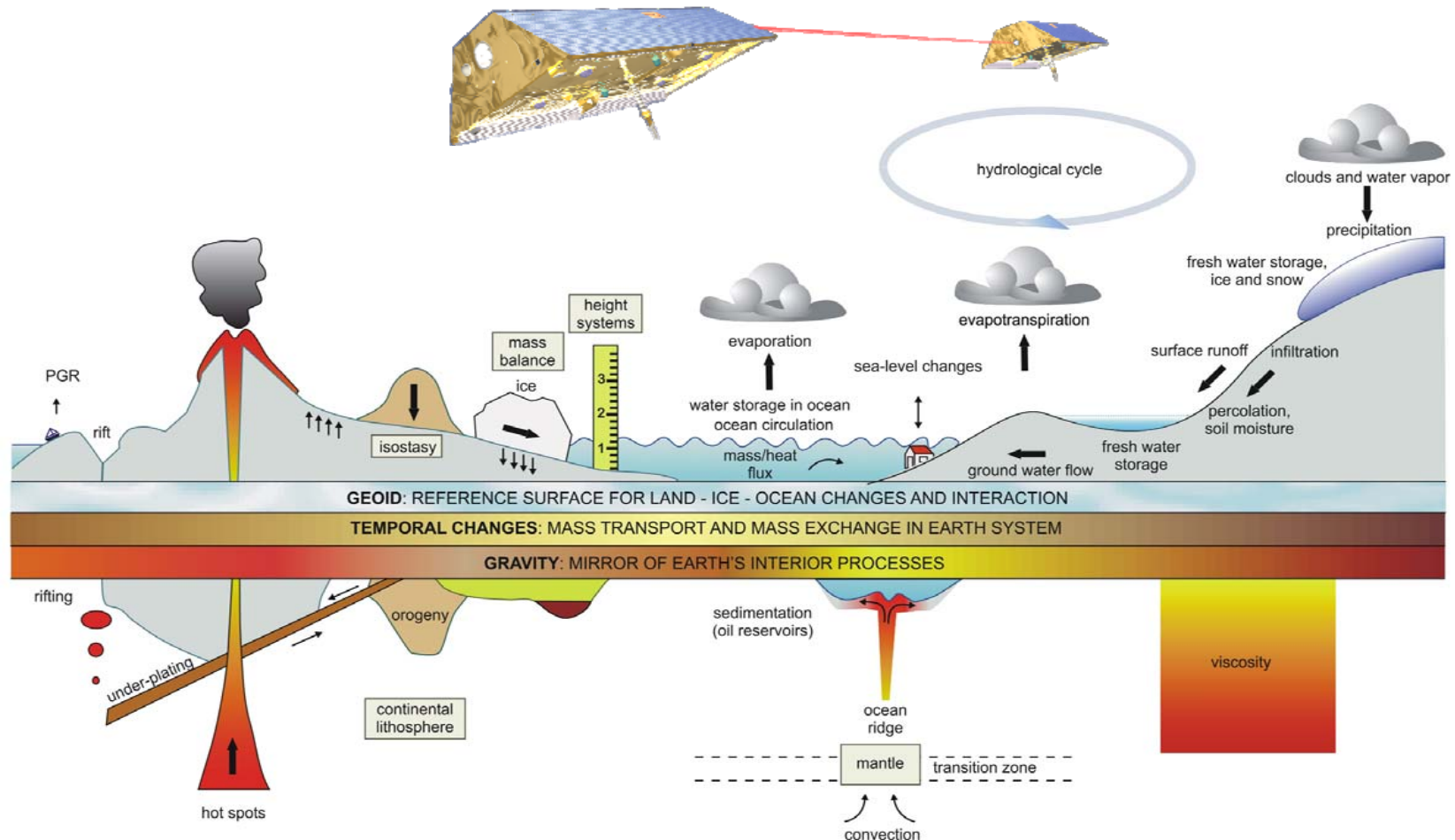
Center for Space Research  
GeoForschungsZentrum  
Jet Propulsion Laboratory

- **Mission Objectives:**

1. Static and time-variable gravity field of the Earth
2. Sounding of the Atmosphere (GPS-Occ.)

# The GRACE Mission

- GRACE represents a quantum leap in gravity recovery enabling observation of mass redistribution processes at unprecedented spatial and temporal scales



from Ilk et al. (2005) Mass Transport and Mass Distribution in the Earth System, 2nd Edition, SPP1257 DFG

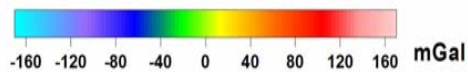
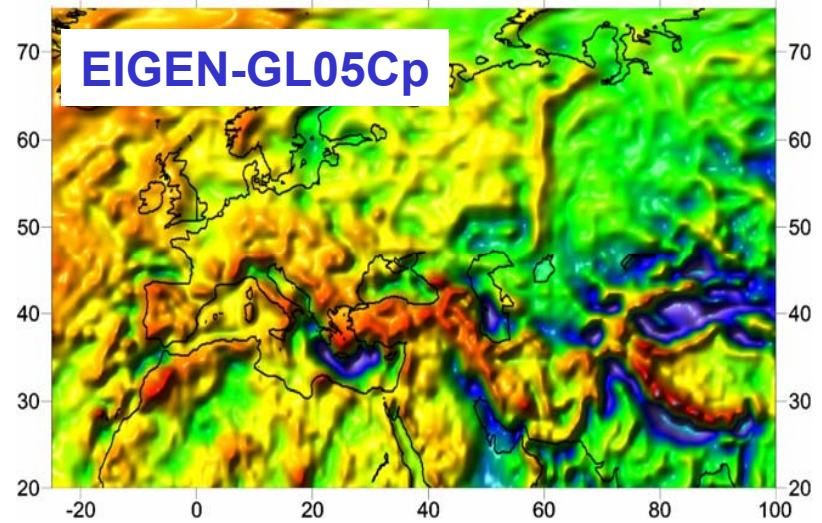
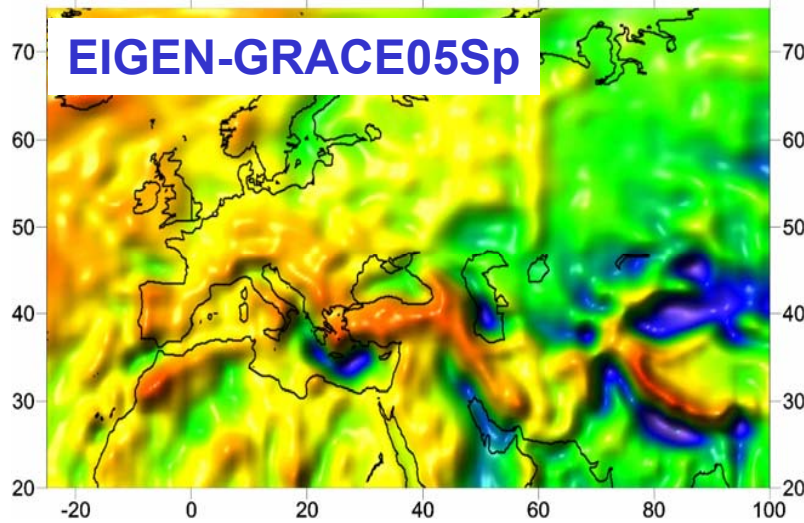
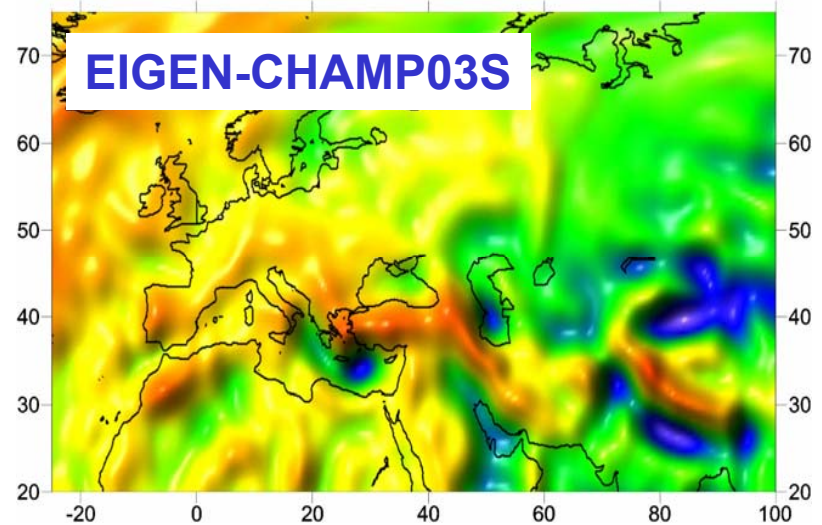
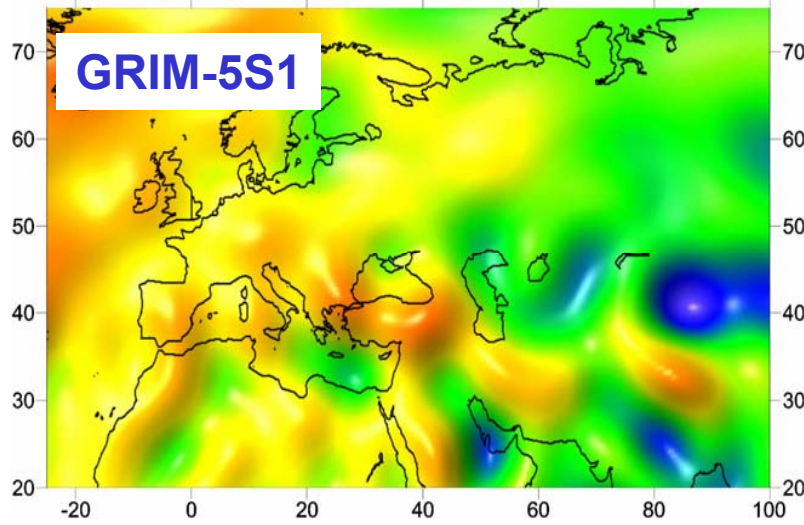


# Overview Results – Static GRACE Gravity Field

- Long-term GRACE-only and ...
  - Spherical harmonics  $N_{\max} \geq 150$
  - Accumulation of  $n > 1$  years of GRACE data
  - $n$  should be integer to reduce seasonal bias through averaging
  - (But secular signals will not be reduced -> Definition of the static field?)
- ... combination gravity field models
  - Use other satellite data (e.g. LAGEOS -> stabilization of C20, in future GOCE?)  
plus
  - terrestrial and altimeter-derived gravity data to increase spatial resolution
  - Spherical harmonics  $N_{\max} = 360$

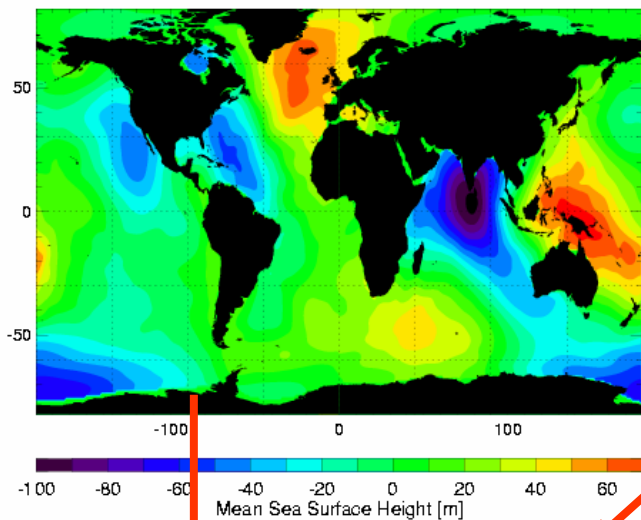


# Overview Results – Static Gravity

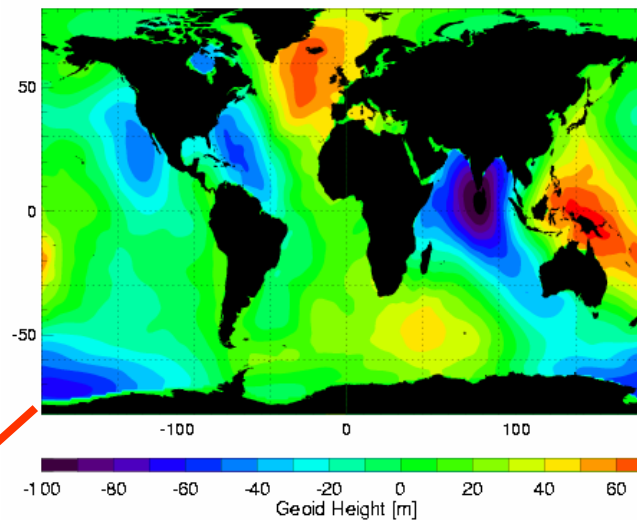


# Dynamic Ocean Topography

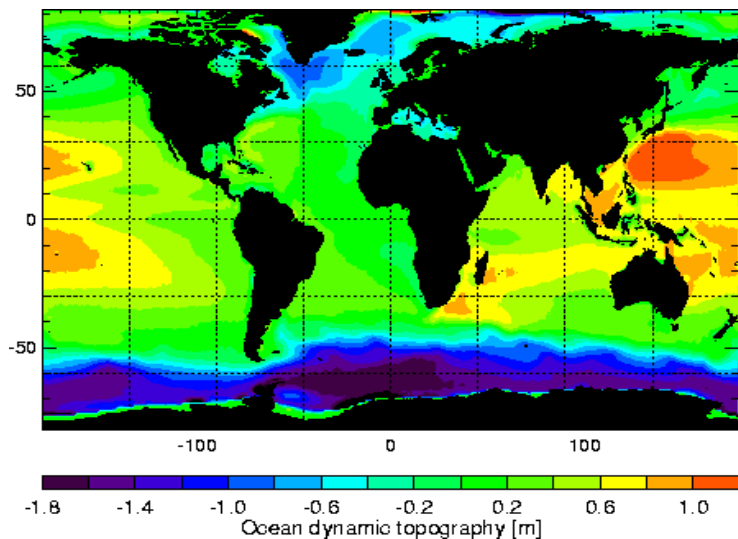
MSSH from Altimetry



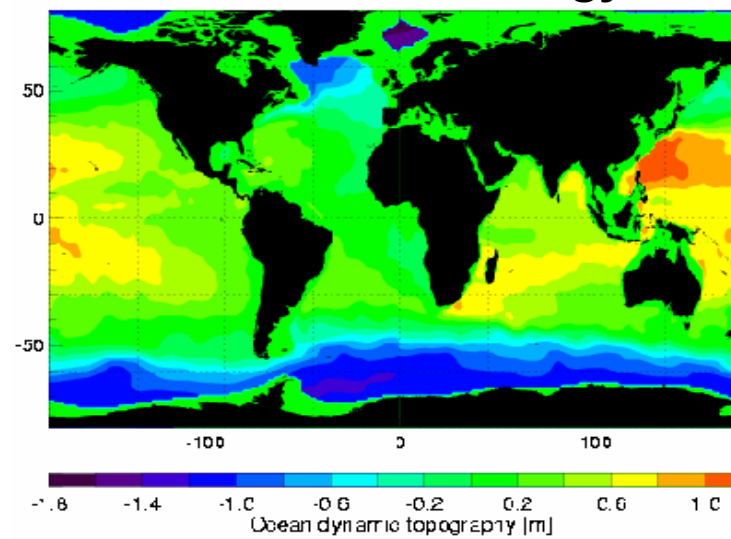
Geoid from EIGEN-GRACE03S (GRACE-only)



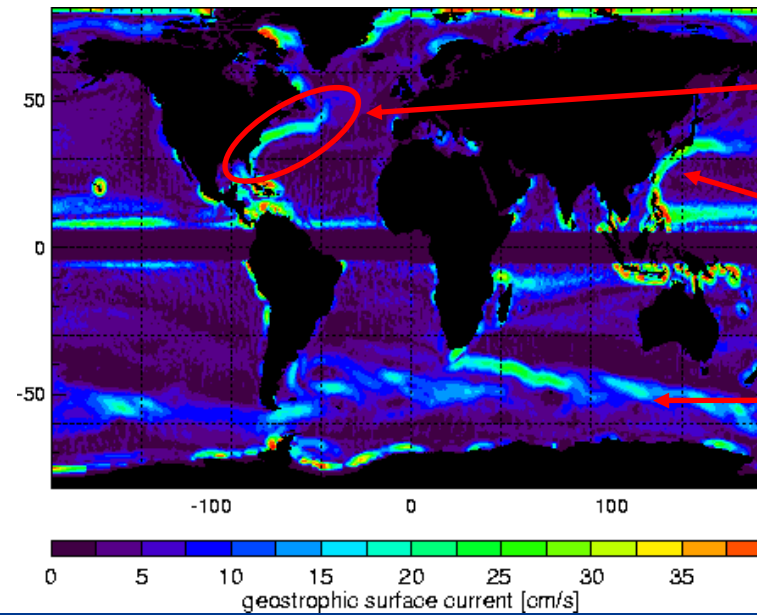
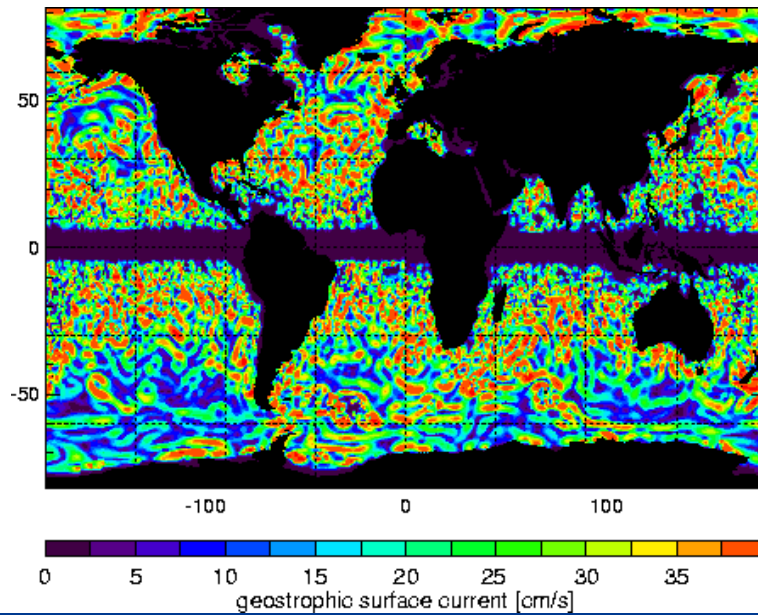
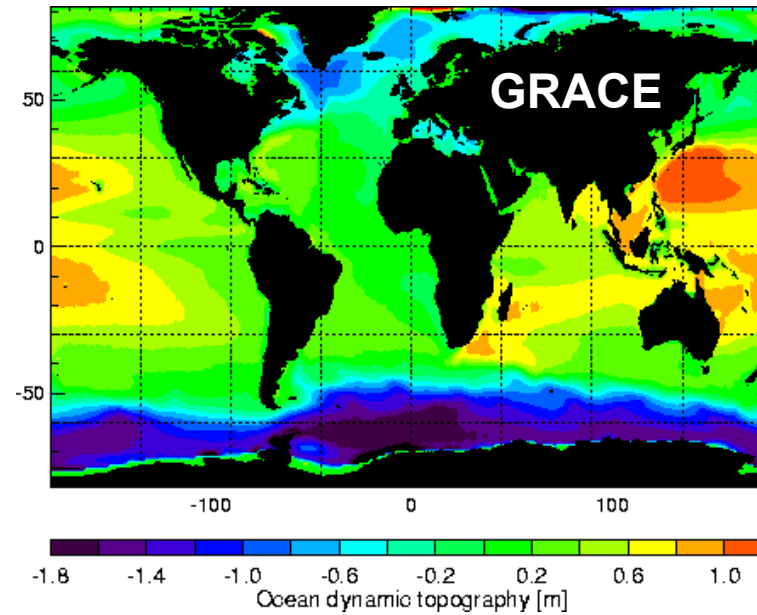
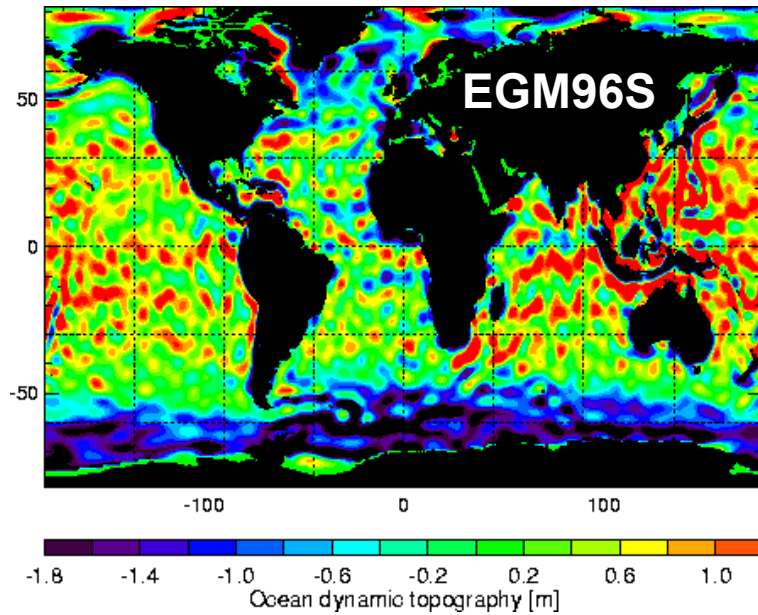
MSSH - EIGEN-GRACE03S



WOCE climatology



# Geostrophic Circulation



Gulf stream

Kuroshio

ACC





# Overview Results – Time-Variable GRACE Gravity

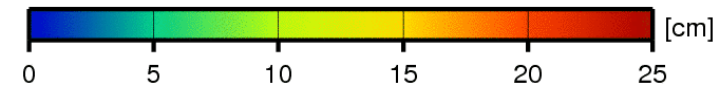
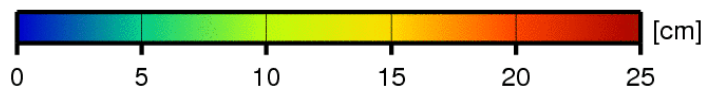
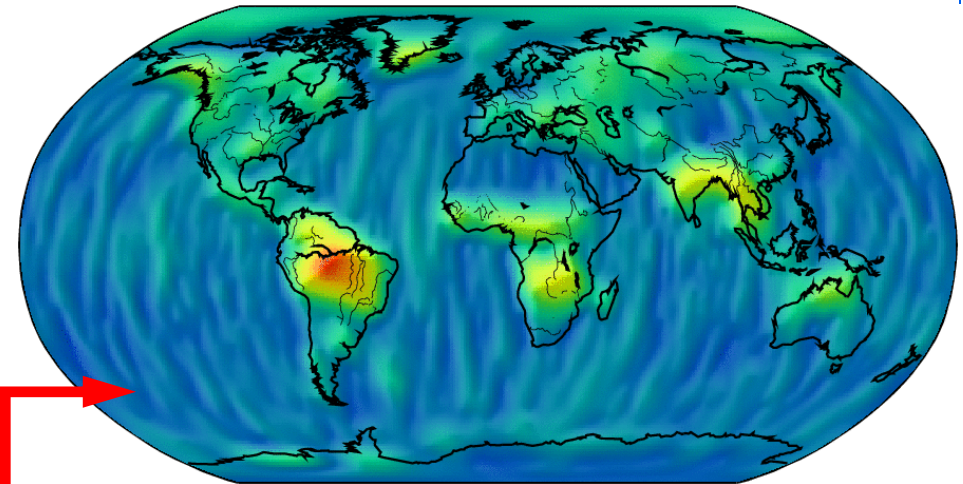
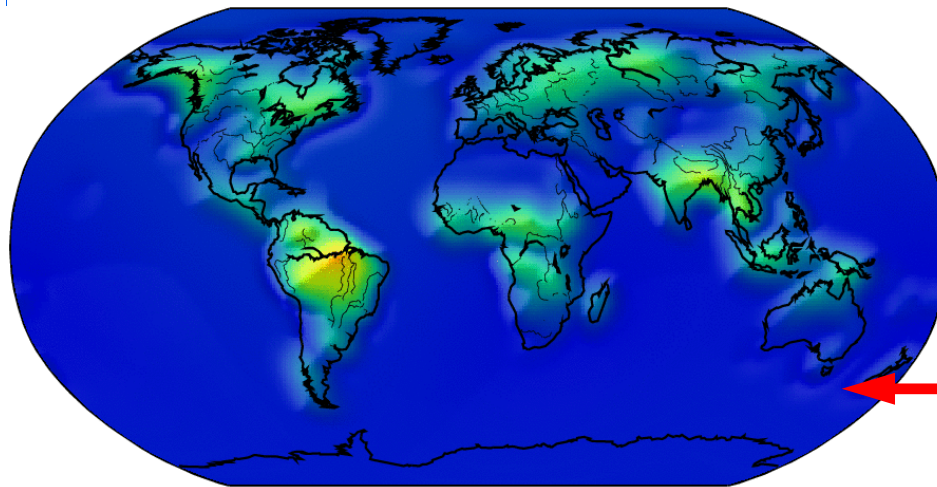
- Provided by GRACE-SDS (CSR, GFZ, JPL) in terms of monthly GRACE-only model series
  - Current release is RL04 (made public begin of 2007)
  - Averages over one calendar month -> temporal evolution discontinuous
  - Global spherical harmonic models (CSR: 60x60, GFZ/JPL: 120x120)
  - Operational data systems providing gravity data at regular intervals (not later than 60 days after L1B data reception)
- From other groups alternative GRACE gravity products exist (not shown)
  - GSFC, GRGS Toulouse, University of Bonn, OSU
  - Different field parametrizations (mascons, radial base functions, ...) and recovery techniques (short arcs, integral approach, energy balance, ...)
  - In general non-operational, irregular delivery

## RMS Surface Mass Anomalies $\sigma(\theta, \lambda, t)$

45 months in 02/2003 – 12/2006  
 (Gaussian averages,  $r = 500$  km)

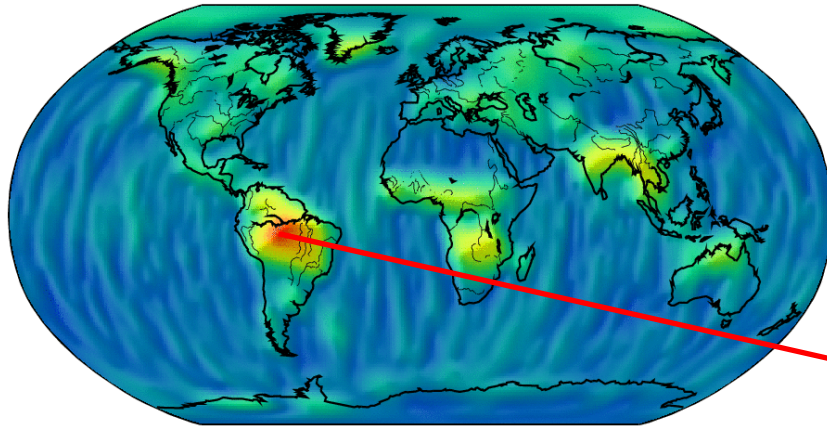
Hydrology Model WGHM\*

GRACE GFZ-RL04

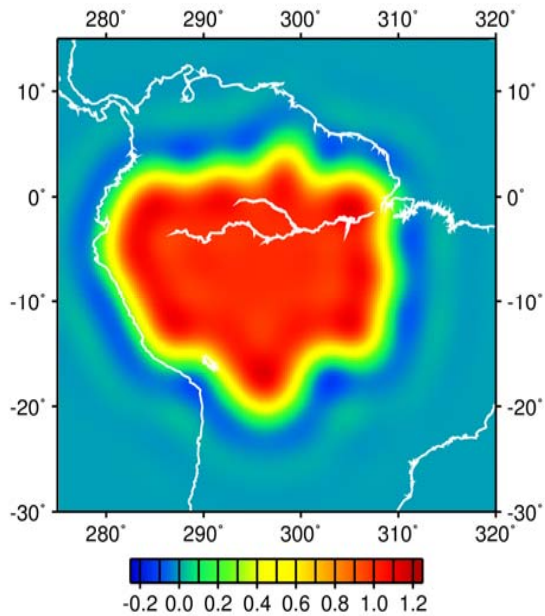


Correlation over land  $\rho > 0.9$

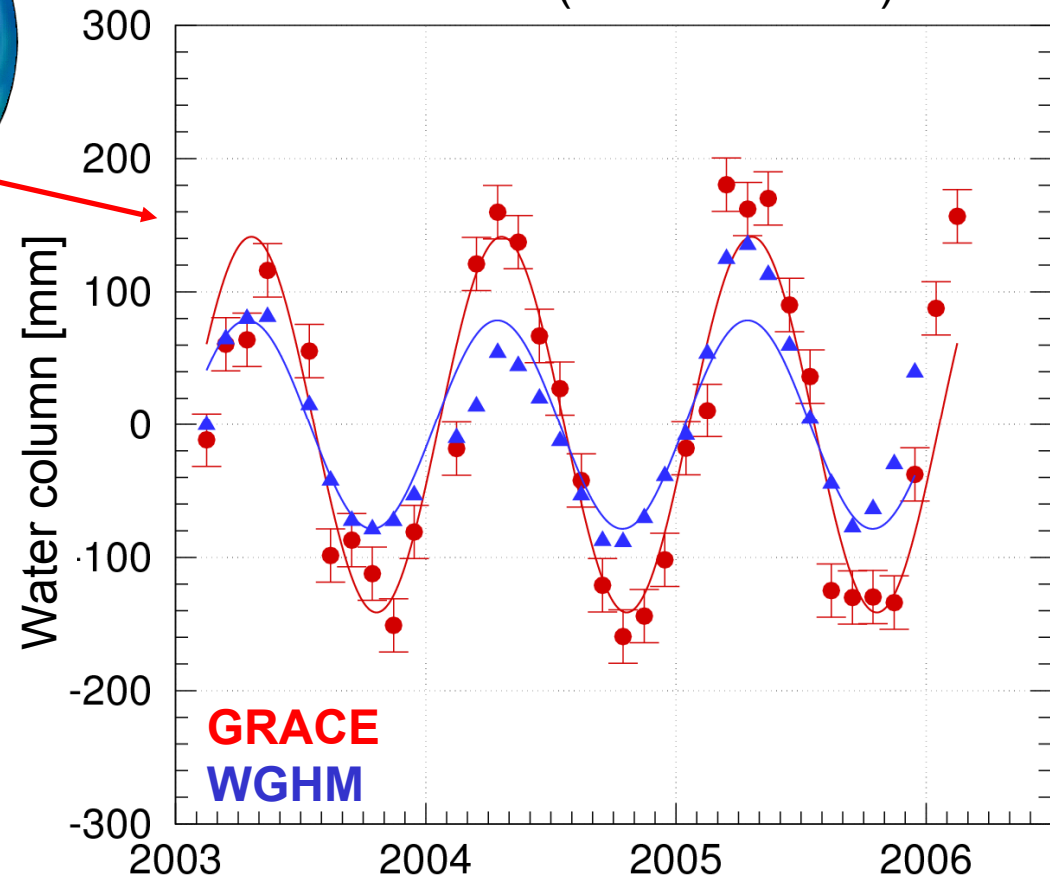
\* WGHM: WaterGAP Hydrology Model (Döll et al. 2003)



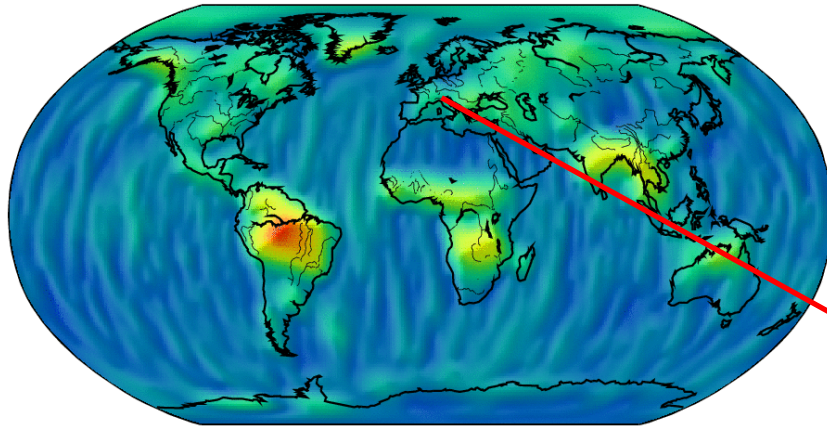
Basin Averages



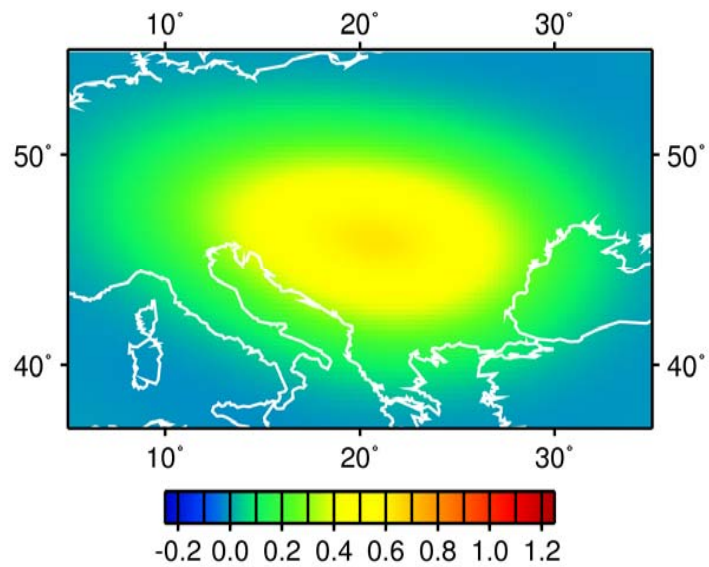
Amazon (5 922 000 km<sup>2</sup>)



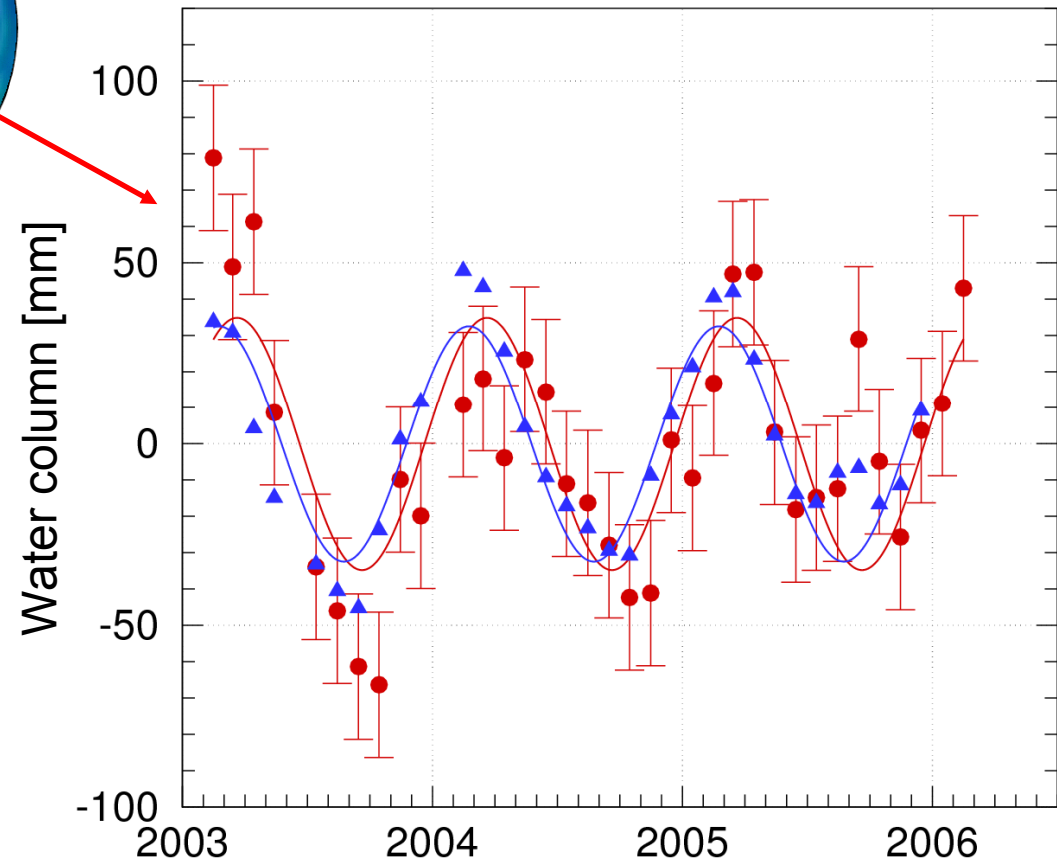
Solid lines = annual sinusoid fitted to data points



Basin Averages



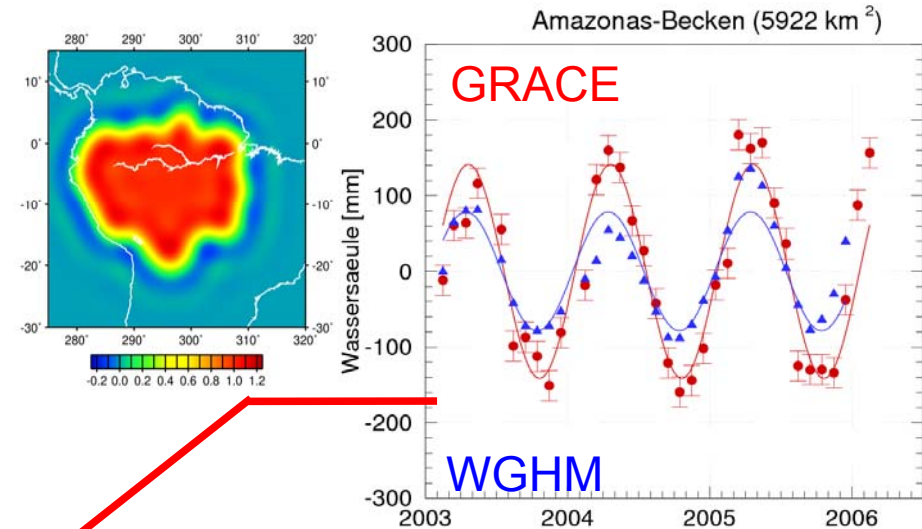
Danube (797 000 km<sup>2</sup>)



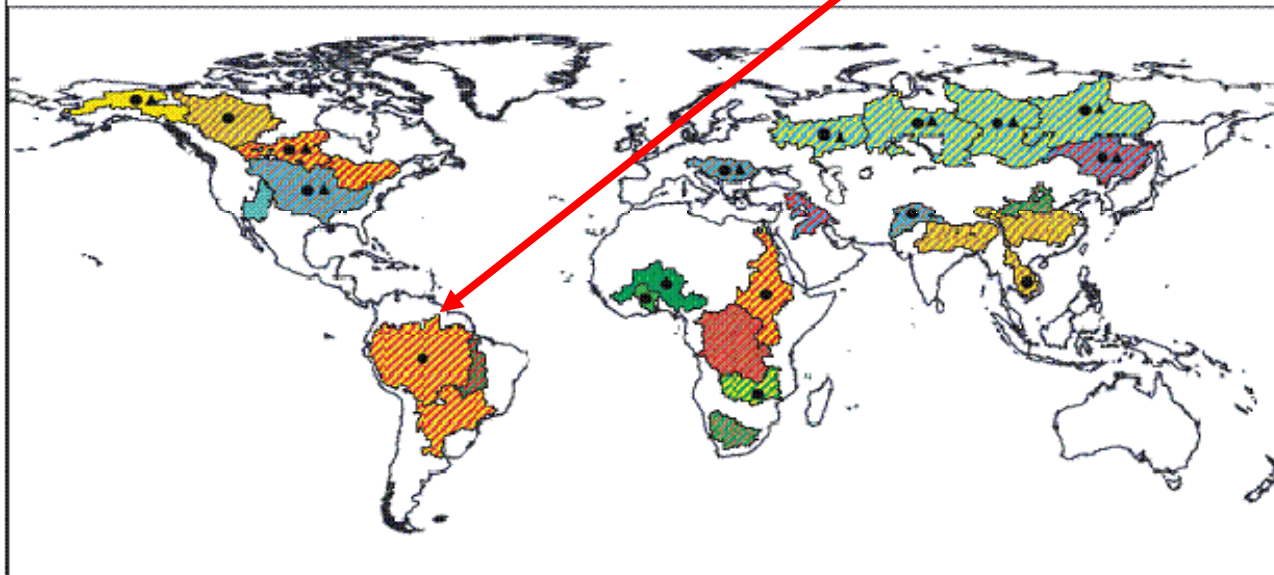
Solid lines = annual sinusoid fitted to data points

## Hydrological aspects:

- Basin-wise calibration required
- Perform sensitivity analysis for hydrological model parameters per basin



Parameter sensitivity for seasonal storage change, WGHM model



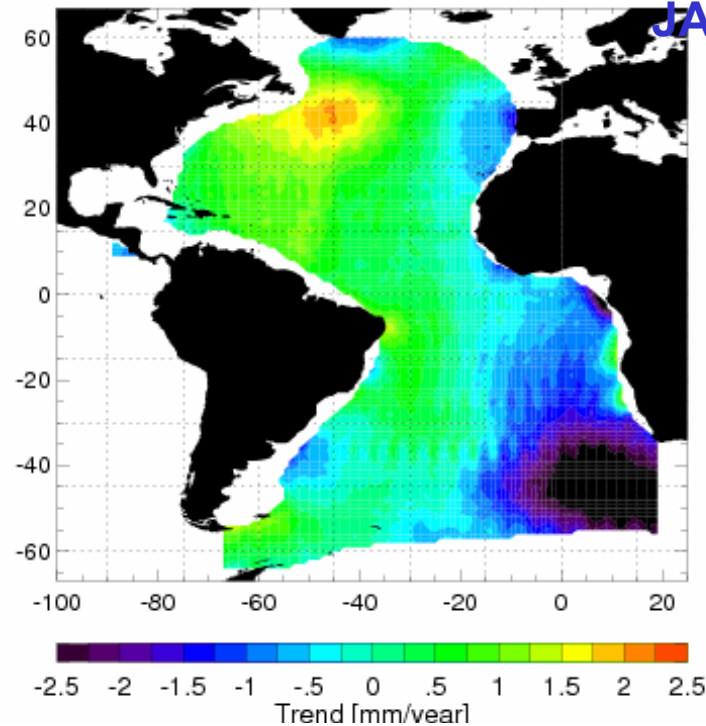
Most sensitive parameters govern processes in the field of

- evapotranspiration / radiation
- soil water
- canopy interception
- snow accumulation / melt
- surface water transport (rivers, lakes, wetlands)

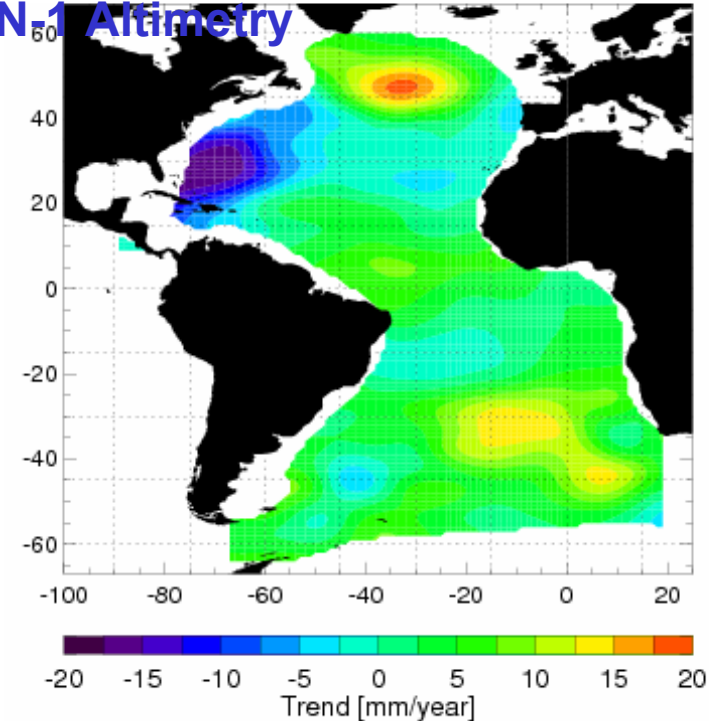
- gamma (original calibration parameter) is highly sensitive
- ▲ precipitation input uncertainties are highly sensitive

## Sea Level Trend Atlantic (08/2002 – 04/2007)

GRACE



JASON-1 Altimetry

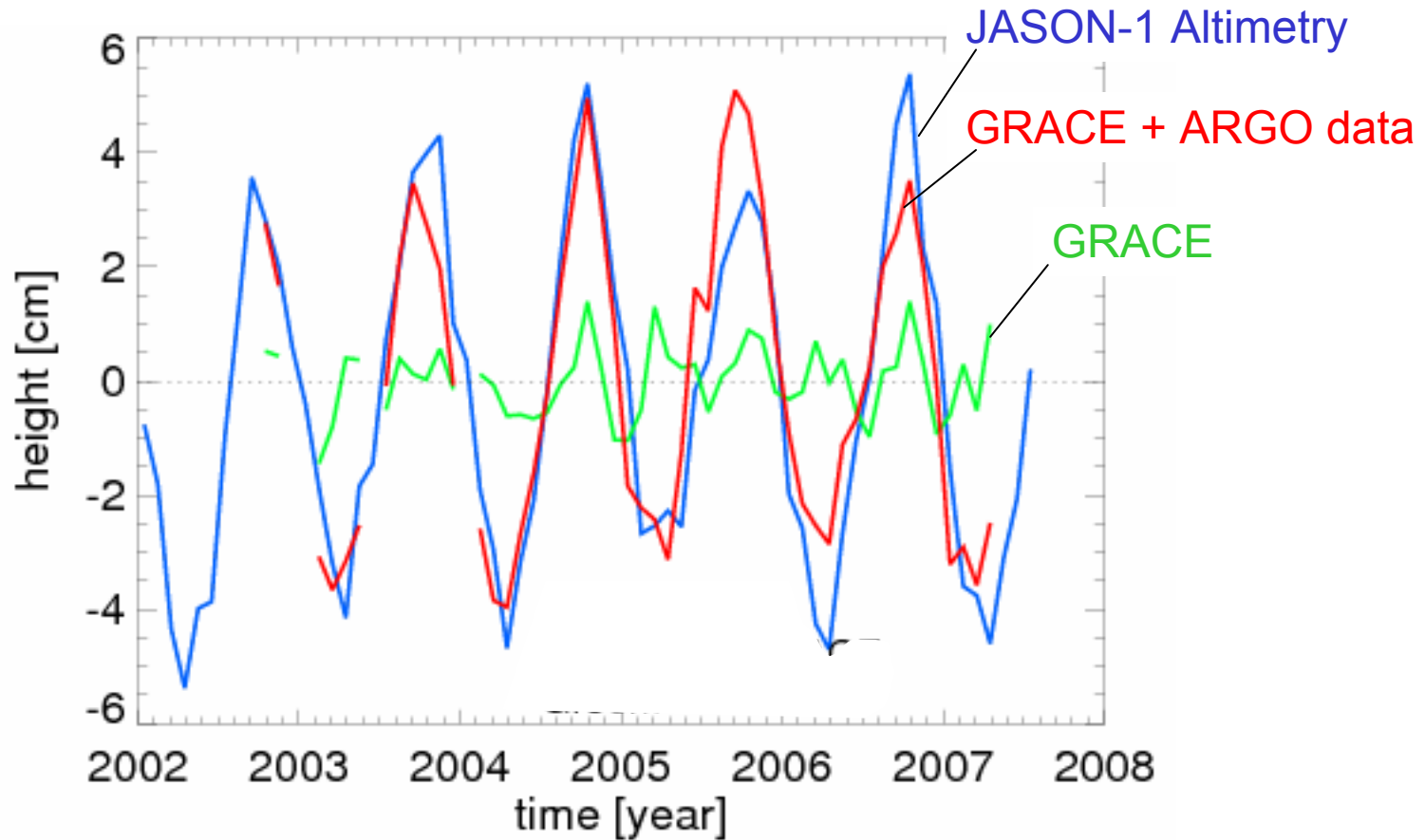


### GRACE:

- Monthly GFZ-RL04 fields 08/2002 – 04/2007 (relative to long-term mean)
- Hamming filter (700 km)
- Short-term ocean mass variability from GAB product re-added
- Geocenter motion from GPS (Heflin, JPL) corrected
- Not yet considered: GIA effects

Courtesy of S. Esselborn, GFZ Potsdam

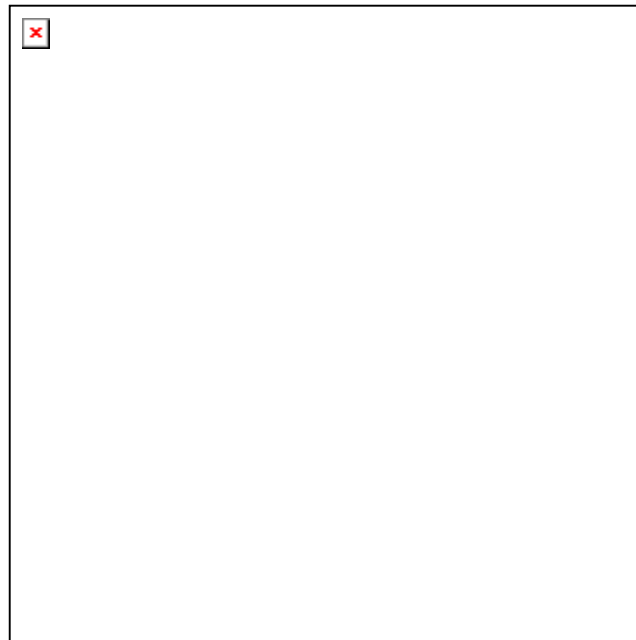
## Mean Sea Level North Atlantic (08/2002 – 04/2007)



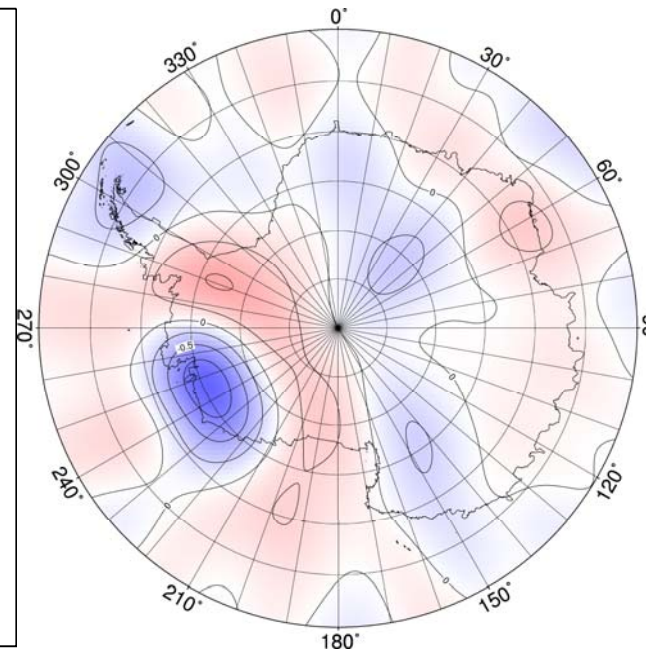
Courtesy of S. Esselborn, GFZ Potsdam

# Ice Mass Loss Over Antarctica

Ice mass balance  
and GIA from models



GRACE  
GFZ RL04 GSM

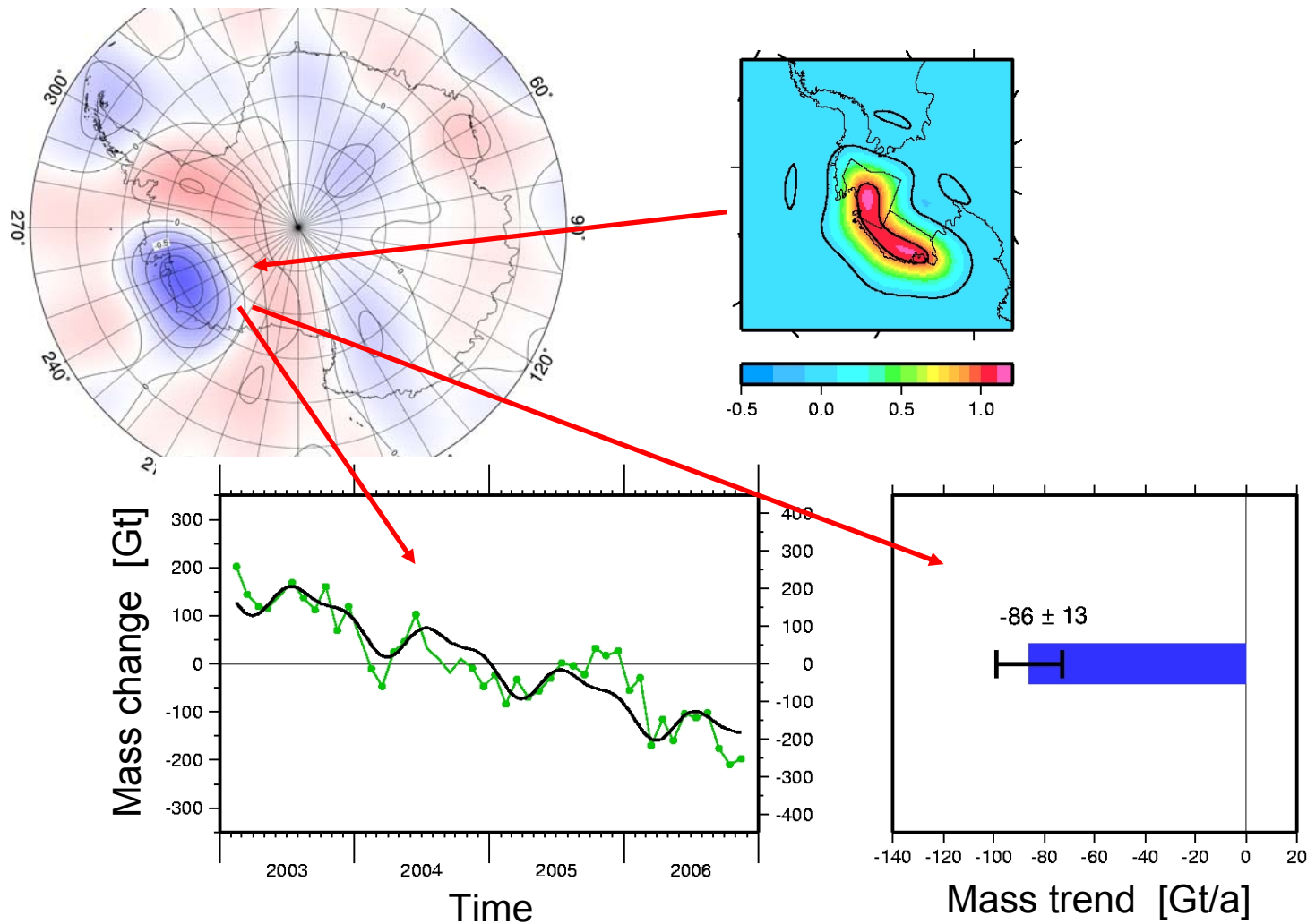


-1.5 -1.0 -0.5 0.0 0.5 1.0 1.5  
Rate of geoid-height change (mm/a)

45 models 02/2003 – 12/2006 (Gaussian averages, filter radius  $r \approx 450$  km)



## Amundsen Sea sector and northwestern Marie Byrd Land



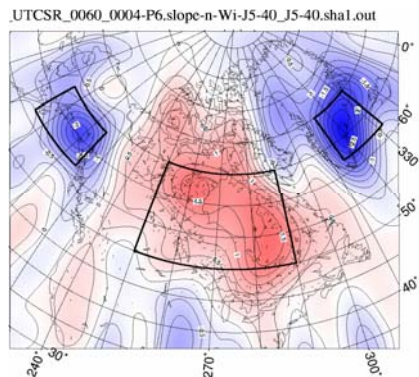
Courtesy of M. Horwath TU Dresden



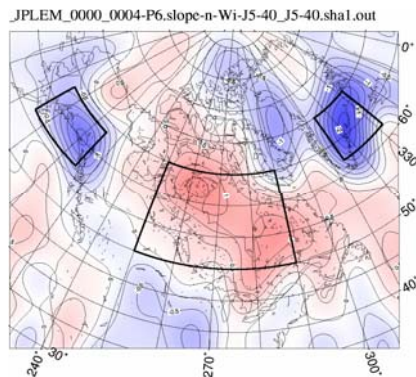
# Trends over Greenland/North America

## GRACE observations

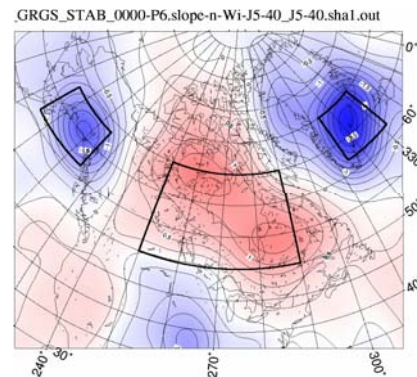
GFZ RL04



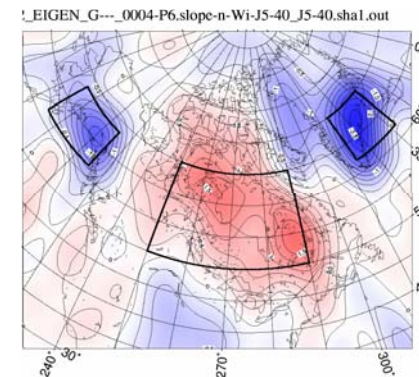
CNES RL01C



JPL RL04

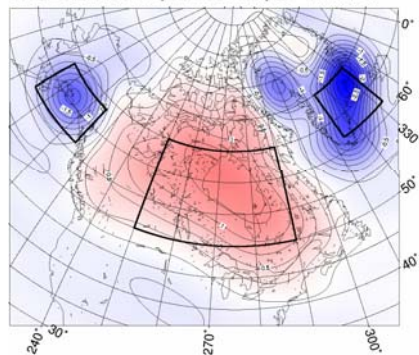


CSR RL04

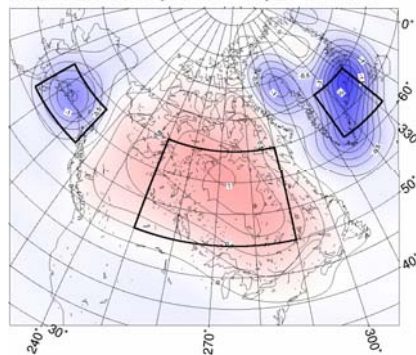


## Adjusted model

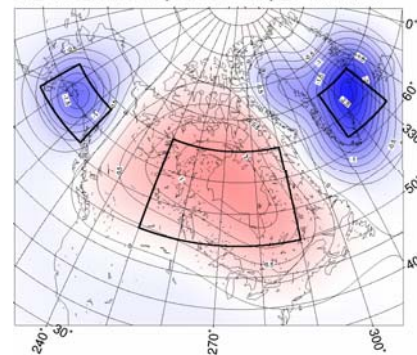
CSR\_0060\_0004-P6.slope-n-Wi-J5-40.opt\_J5-40.sha1.out



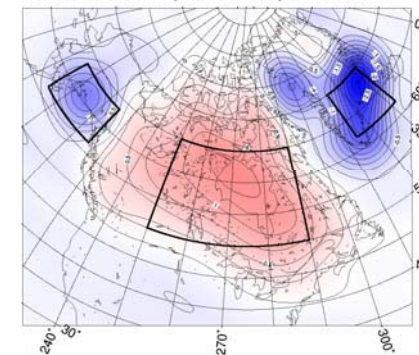
EM\_0000\_0004-P6.slope-n-Wi-J5-40.opt\_J5-40.sha1.out



GS\_STAB\_0000-P6.slope-n-Wi-J5-40.opt\_J5-40.sha1.out



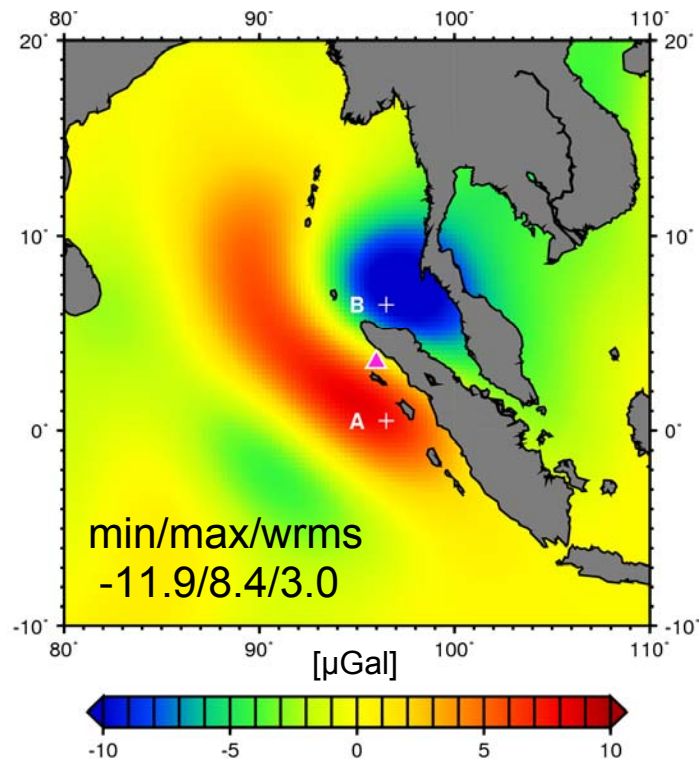
GEN\_G---\_0004-P6.slope-n-Wi-J5-40.opt\_J5-40.sha1.out



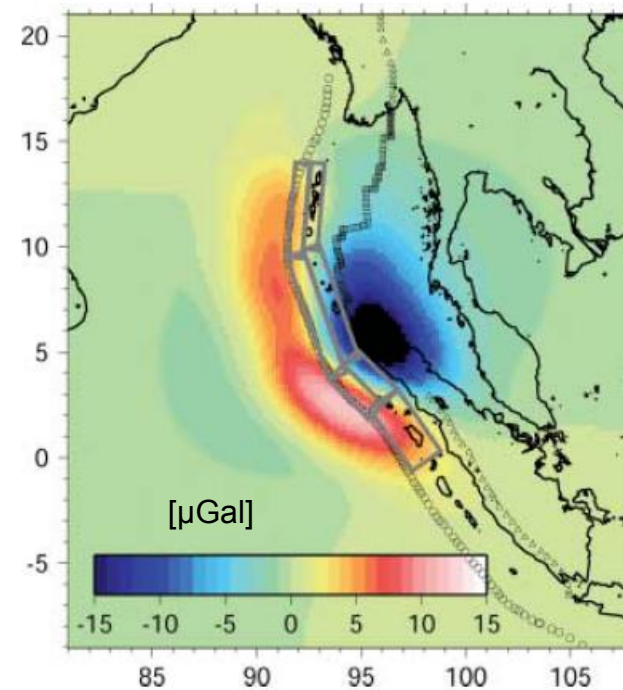
Courtesy of I. Sasgen, GFZ Potsdam

## GFZ-RL04

Differences of averages<sup>\*)</sup> from 24 monthly models *before* and 29 monthly models *after* the EQ (residual to mean from 53 months)



Predicted coseismic signal (Han et al. 2006)

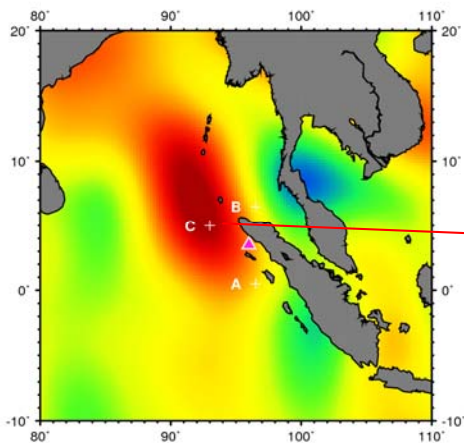


<sup>\*)</sup> De-stripping using decorrelation method by J. Kusche, Journal of Geodesy, 2007

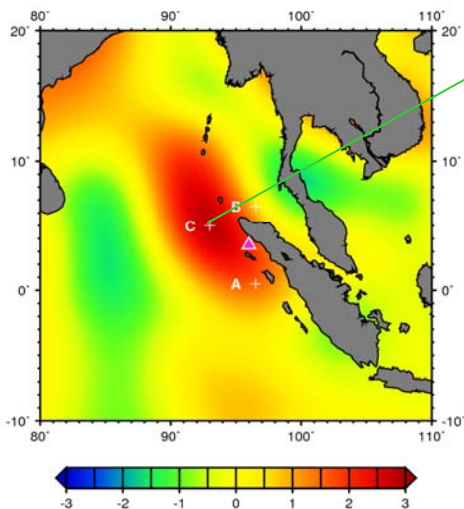


# Local Non-Seasonal Gravity Change

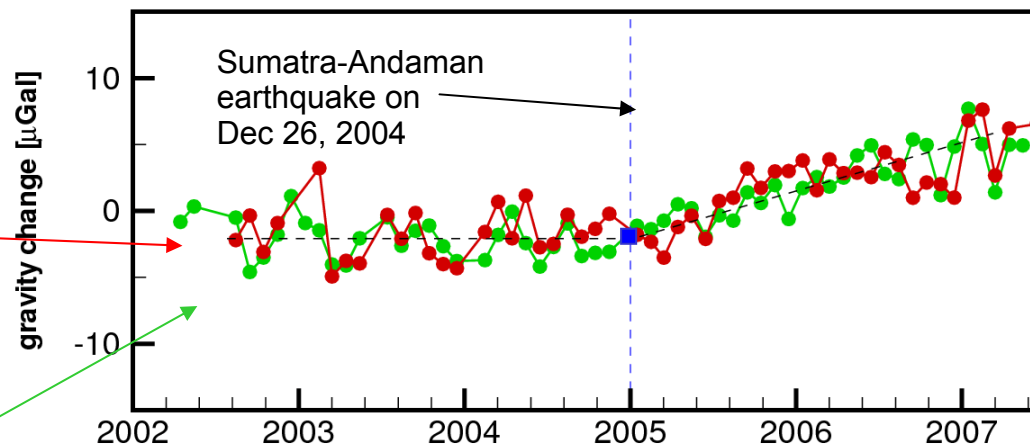
GFZ-RL04



CSR-RL04



Point C (5° N, 93° E)



Data	Trend before EQ	
Trend after EQ		
GFZ-RL04	$0.4 \pm 0.6$	$3.1 \pm 0.5$
CSR-RL04	$-0.6 \pm 0.4$	$2.9 \pm 0.4$
Model	-	tbd

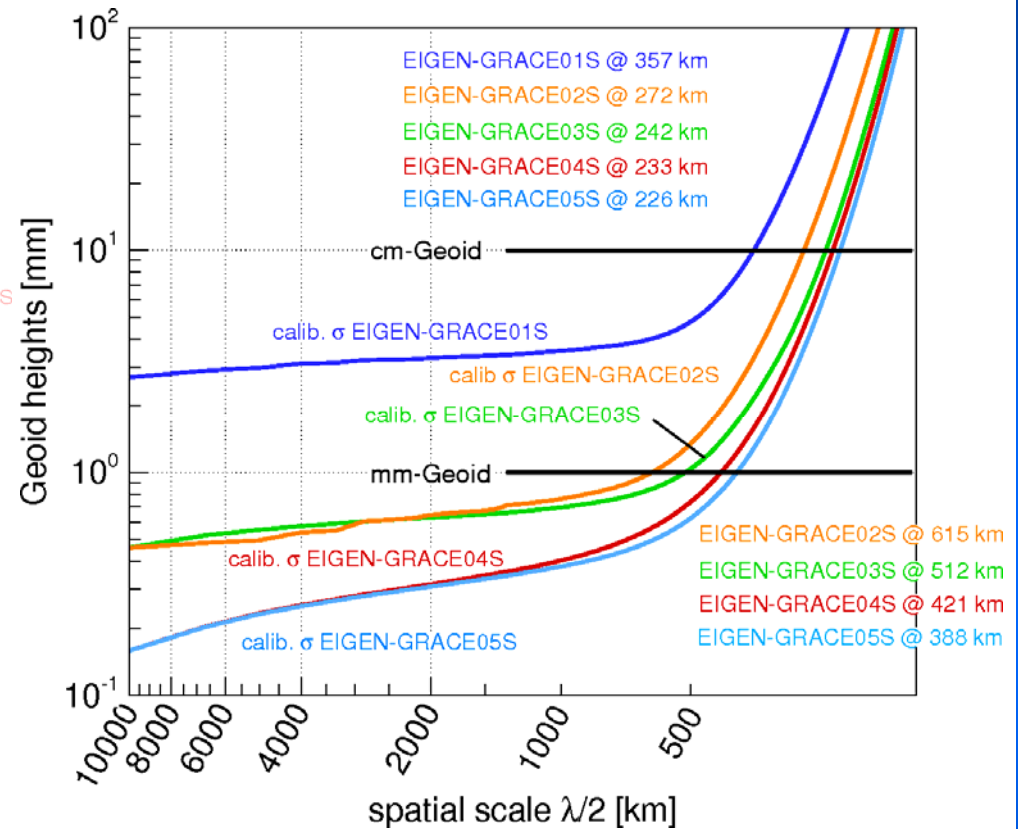
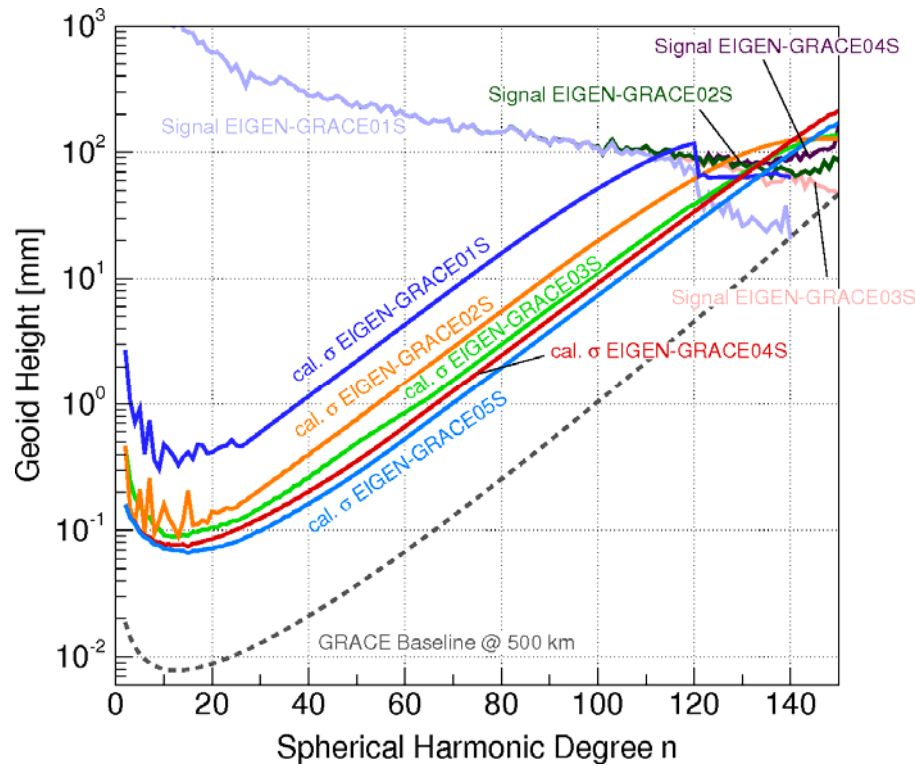
Units: [ $\mu\text{Gal}/\text{a}$ ], formal errors from fit

Variability at location A, B **after** fitting annual and semiannual terms for each interval (2002-2004 w/o 12/2004) and (2005-2007) separately

■ December 2004 excluded



# Evolution of EIGEN-GRACE Models



## Error level *static* GRACE-only fields:

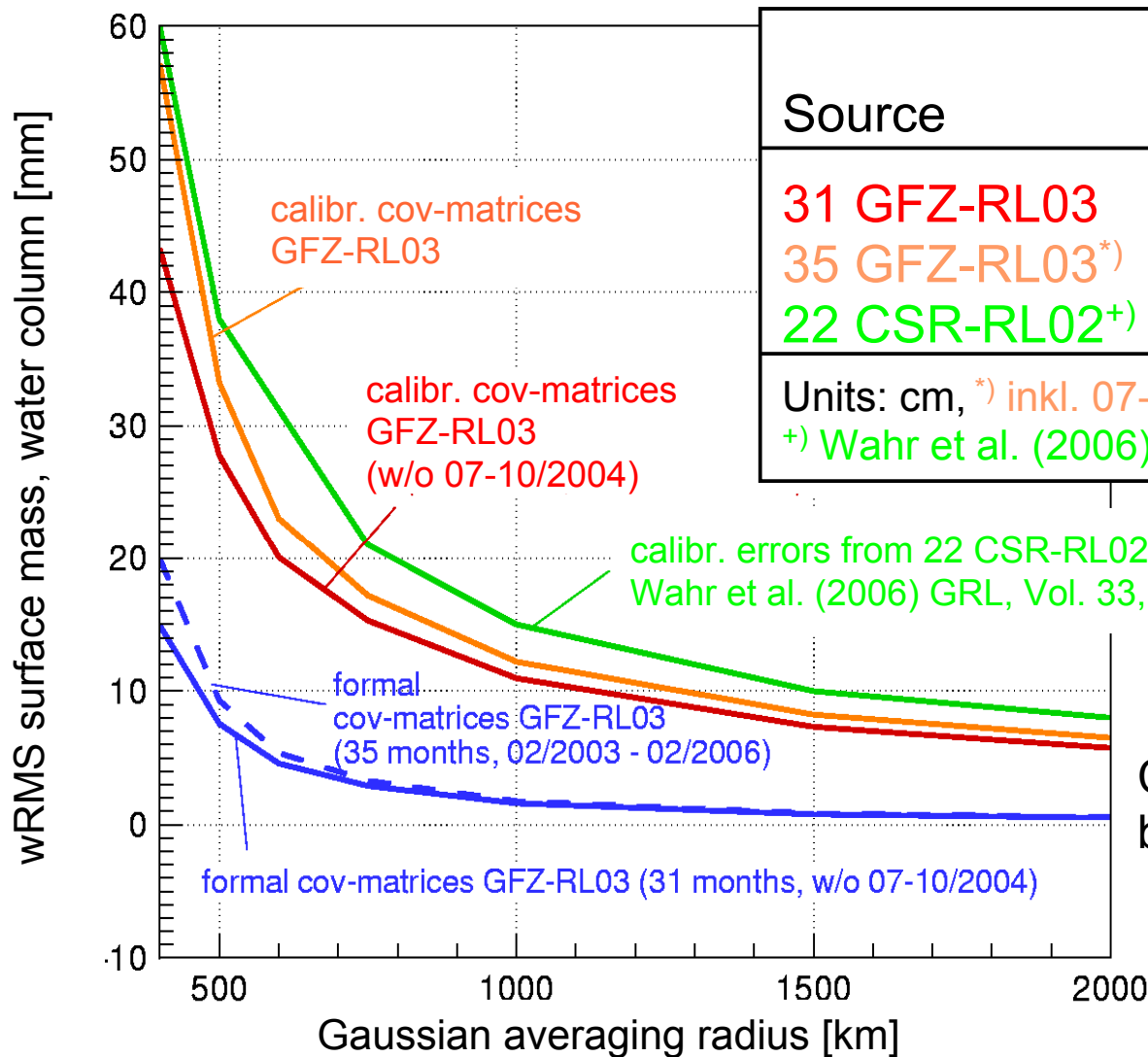
EIGEN-GRACE01S (=RL00)	≈ 50 x baseline
EIGEN-GRACE02S (=RL01)	≈ 20 x baseline
EIGEN-GRACE03S (=RL02)	≈ 12.5 x baseline
EIGEN-GRACE04S (=RL03)	≈ 10 x baseline
EIGEN-GRACE05S (=RL04)	≈ 7.5 x baseline

## Error level *monthly* GRACE-only models:

EIGEN-GRACE01S (=RL00)	no monthly field series
EIGEN-GRACE02S (=RL01)	≈ 40 x baseline
EIGEN-GRACE03S (=RL02)	≈ 25 x baseline
EIGEN-GRACE04S (=RL03)	≈ 17.5 x baseline
EIGEN-GRACE05S (=RL04)	≈ 15 x baseline



# Evolution of EIGEN-GRACE Models



Source	filter radii [km]			
	1500	1000	750	500
31 GFZ-RL03	0.7	1.1	1.5	2.7
35 GFZ-RL03 <sup>*)</sup>	0.8	1.3	1.8	3.5
22 CSR-RL02 <sup>+) )</sup>	1.0	1.5	2.1	3.9

Units: cm, <sup>\*)</sup> inkl. 07-10/2004  
<sup>+) )</sup> Wahr et al. (2006) GRL, Vol. 33, L06401

Calibration of GRACE errors by means of internal comparisons



# Current Questions (1)

- Assessment and treatment of correlated errors (“stripes”) in GRACE gravity products
  - Determination of external accuracy of GRACE gravity products
  - Post-processing of gravity models required. Any filtering technique causes biases in amplitudes/phases of GRACE-products
- Refinements in data processing to reduce/avoid aliasing effects
  - Improved background models (e.g. ocean tides, short-term mass variations,...)
  - Alternative recovery methods, refined parametrizations (gravity field, orbit, instruments, ...)
  - To what extent one can “eliminate” a GRACE background model (e.g. ocean) in post-processing?



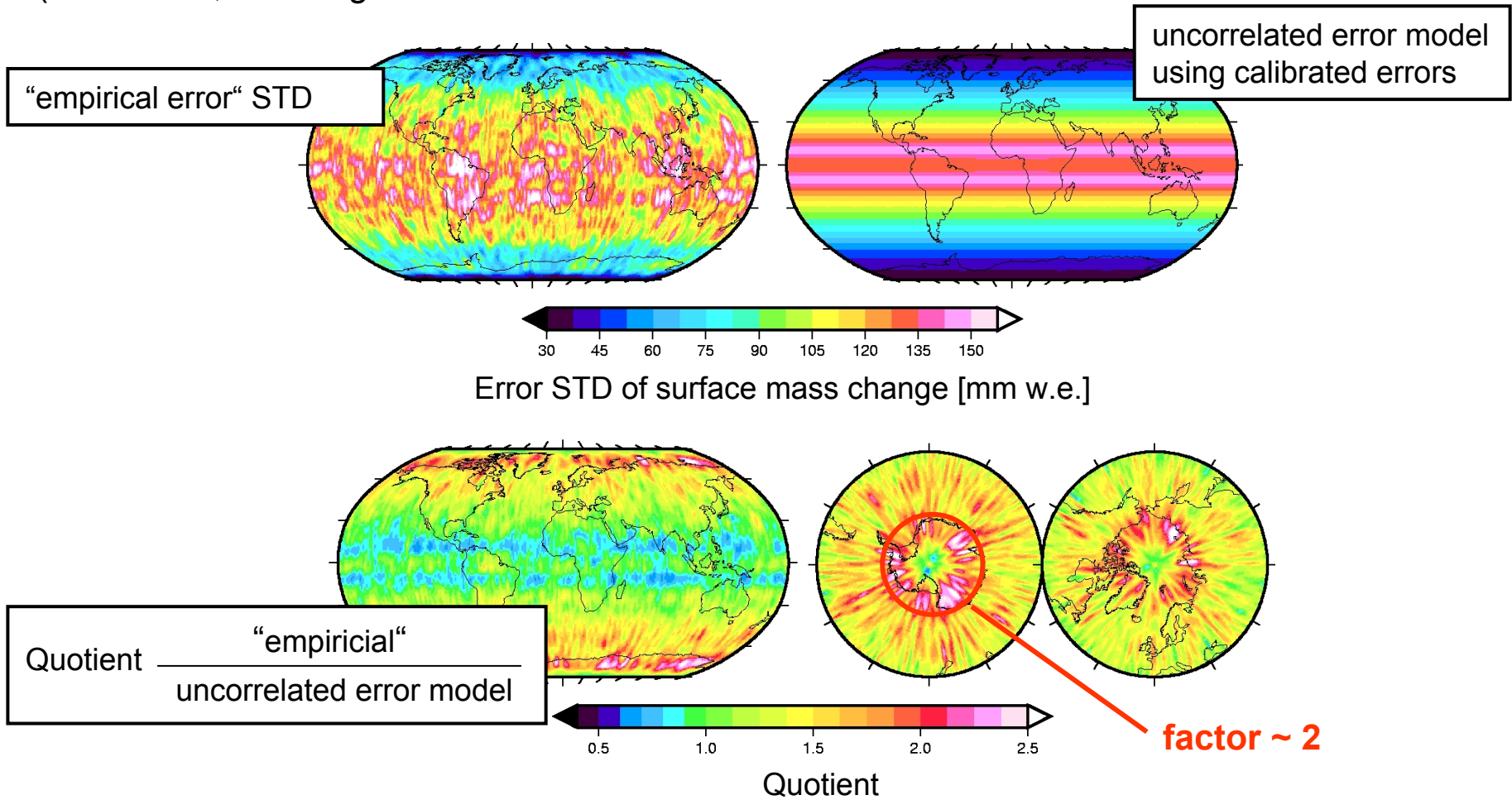
## Current Questions (2)

- Additional, in general, but also relevant for GGOS:
  - Consistency of GRACE gravity products and GRACE-derived surface mass changes, products from geometrical techniques, Earth rotation parameters, results from altimetry?
  - Current GRACE data products sufficient for GGOS? With respect to consistency but also spatial and temporal resolution?
  - Continuation of GRACE-type data as an operational part of GGOS (GRACE Follow-On)?



# Impact of Correlations

Empirical errors vs. propagated calibrated errors [Horwath and Dietrich, GRL, 2006]  
 (GFZ RL04, 300km gaussian smoothed surface mass



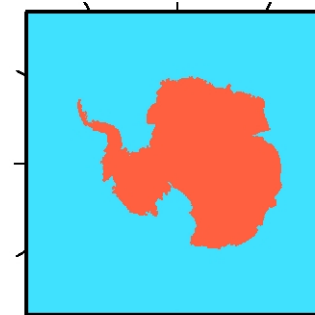
Integrated mass signal

$$\Delta m(t) = \iint_{\Omega} \vartheta(\Omega) \Delta \sigma^{\text{true}}(\Omega, t) a^2 d\Omega$$

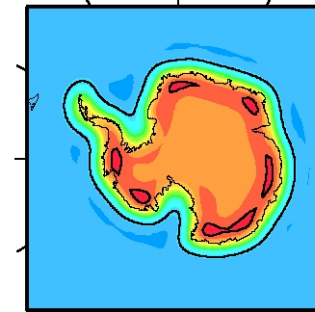
Surface mass density

GRACE mass estimate

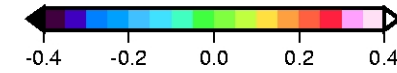
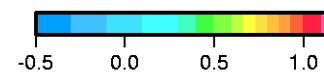
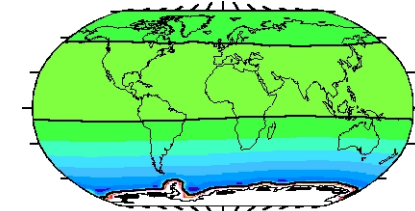
$$\begin{aligned} \widehat{\Delta m}(t) &= \sum_{\{nm\}_{\text{sat}}} \alpha_{nm} \Delta c_{nm}^{\text{sat}}(t) \\ &= \iint_{\Omega} \eta(\Omega) \Delta \sigma^{\text{sat}}(\Omega, t) a^2 d\Omega \end{aligned}$$



region function  $\vartheta$



weight function  $\eta$



Errors

- $\eta \neq \vartheta$  → Leakage effect (induced by any uncorrected  $\Delta \sigma$ )
- $\Delta \sigma^{\text{sat}} \neq \Delta \sigma^{\text{true}}$  → GRACE error effect
- $\Delta \sigma^{\text{true}} = \Delta \sigma_{\text{ice}} + \Delta \sigma_{\text{GIA}} + \Delta \sigma_{\text{others}}$  → Errors of superimposed signal correction

All error effects depend on  $\alpha_{nm}$  or, respectively,  $\eta$



## **JIGOG = Joint Inversion of GPS site displacements, ocean bottom pressure models and GRACE gravimetry**

- **DFG-project (SPP 1257 “Mass Transport”):**

- Partners



- **Data:**

- GRACE: GFZ-RL04 gravity models, weekly resolution
- GPS station displacements from IGS
- Ocean Bottom Pressure (OBP) data from AWI FESOM model

- **Key features:**

- Global time variable mass transport (ocean/atmosphere, hydrology)
- Mass consistent hybrid product from three data sets (optimal in least square sense)
- Estimate geocenter motion simultaneously
- Low spatial resolution

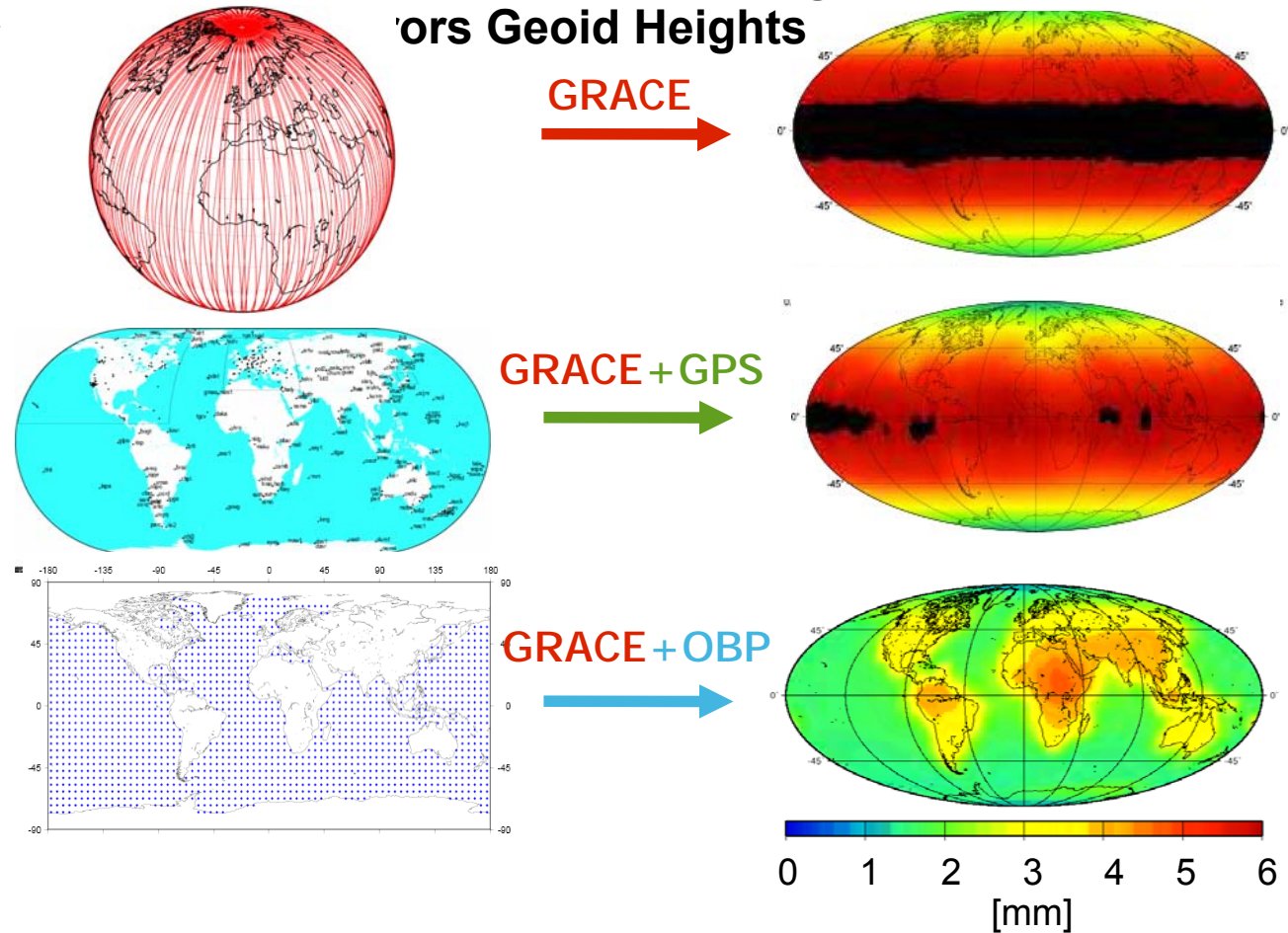
## Stabilization

### Data

All data sets with weekly resolution

- GRACE GFZ-RL04
- IGS site displacements
- OBP from FESOM

### Spatial Coverage for Geoid Heights



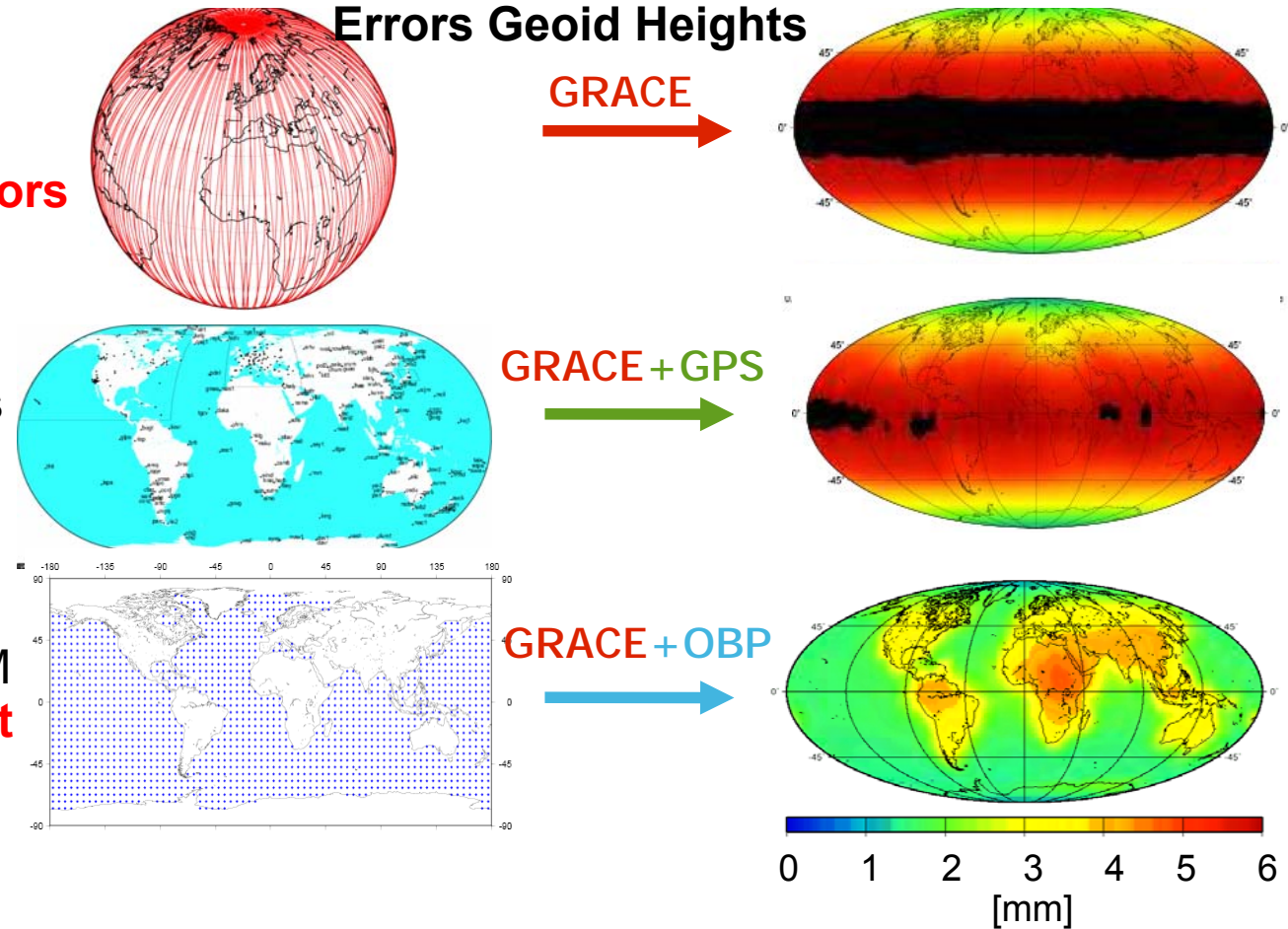
## Stabilization

### Data

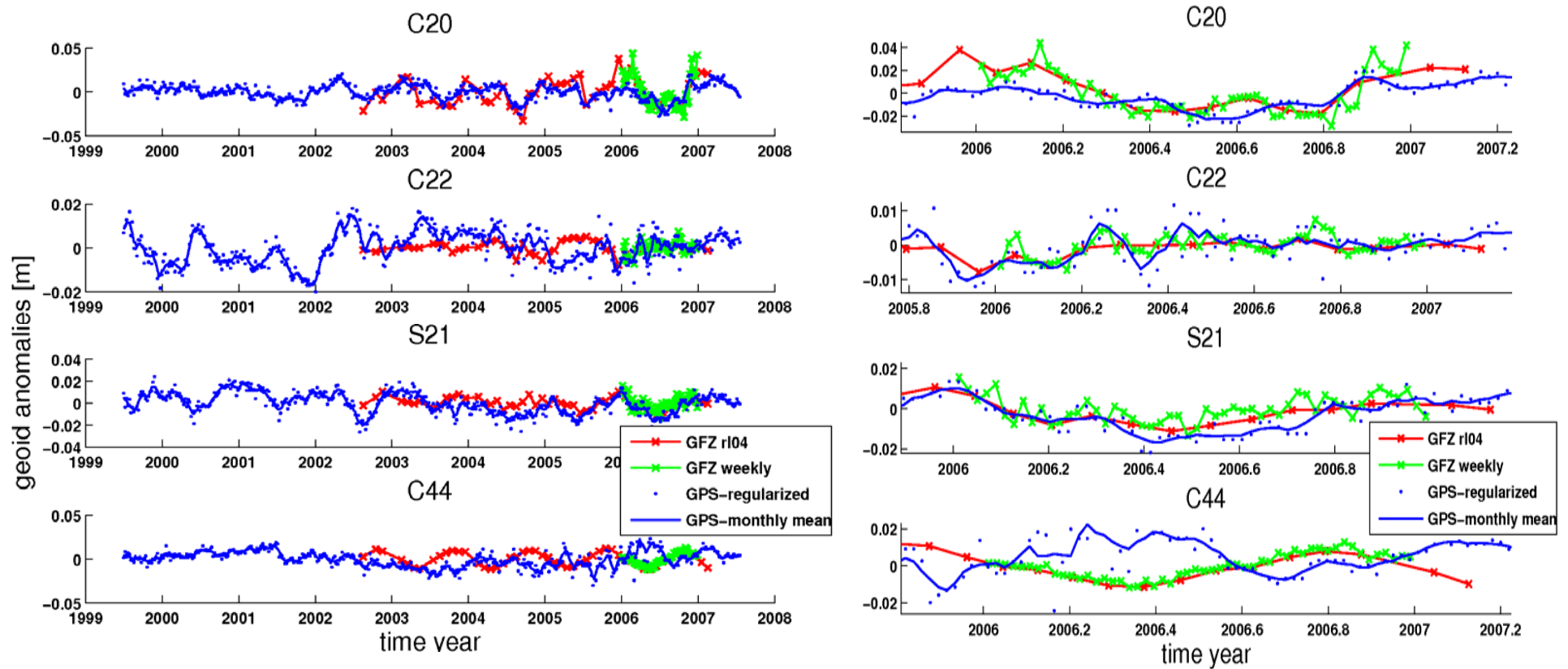
#### Main Challenges

- GRACE GFZ-RL04  
**Systematic model errors**
- IGS site displacements  
**Data distribution**
- OBP from AWI FESOM  
**Accuracy assessment**

### Spatial Coverage Errors Geoid Heights



## Comparison low degree/order coefficients



- Good agreement for C20 and S21
- Weekly GFZ-RL04 models realistic but noisier

Courtesy of R. Rietbroek, GFZ Potsdam



# Conclusions

- GRACE-based long-term and time-variable gravity models provide important novel input on the spatio-temporal mass distribution within the Earth system for applications in oceanography, hydrology, glaciology, geophysics (Earth's interior, seismic deformation due to large Earthquakes)
  - Quality of GRACE static and time-variable gravity field models has been improved from release to release (but baseline still not reached)
- However, to exploit full benefit from GRACE for derivation of mass transport signals complementary data and observation techniques must be used which will also be mandatory for a usage for GGOS. Efforts will be directed towards:
  - Accuracy assessment
  - Reduction of correlated errors
  - Consistency of gravity and geometry in a common reference frame
  - GRACE-based surface mass estimates to be used for loading computations

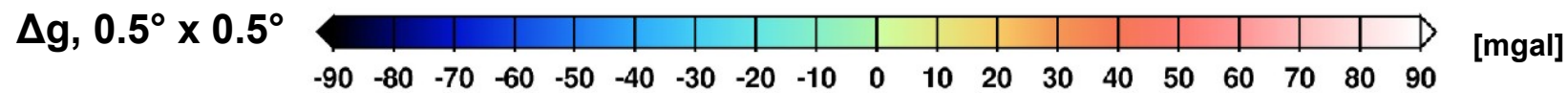
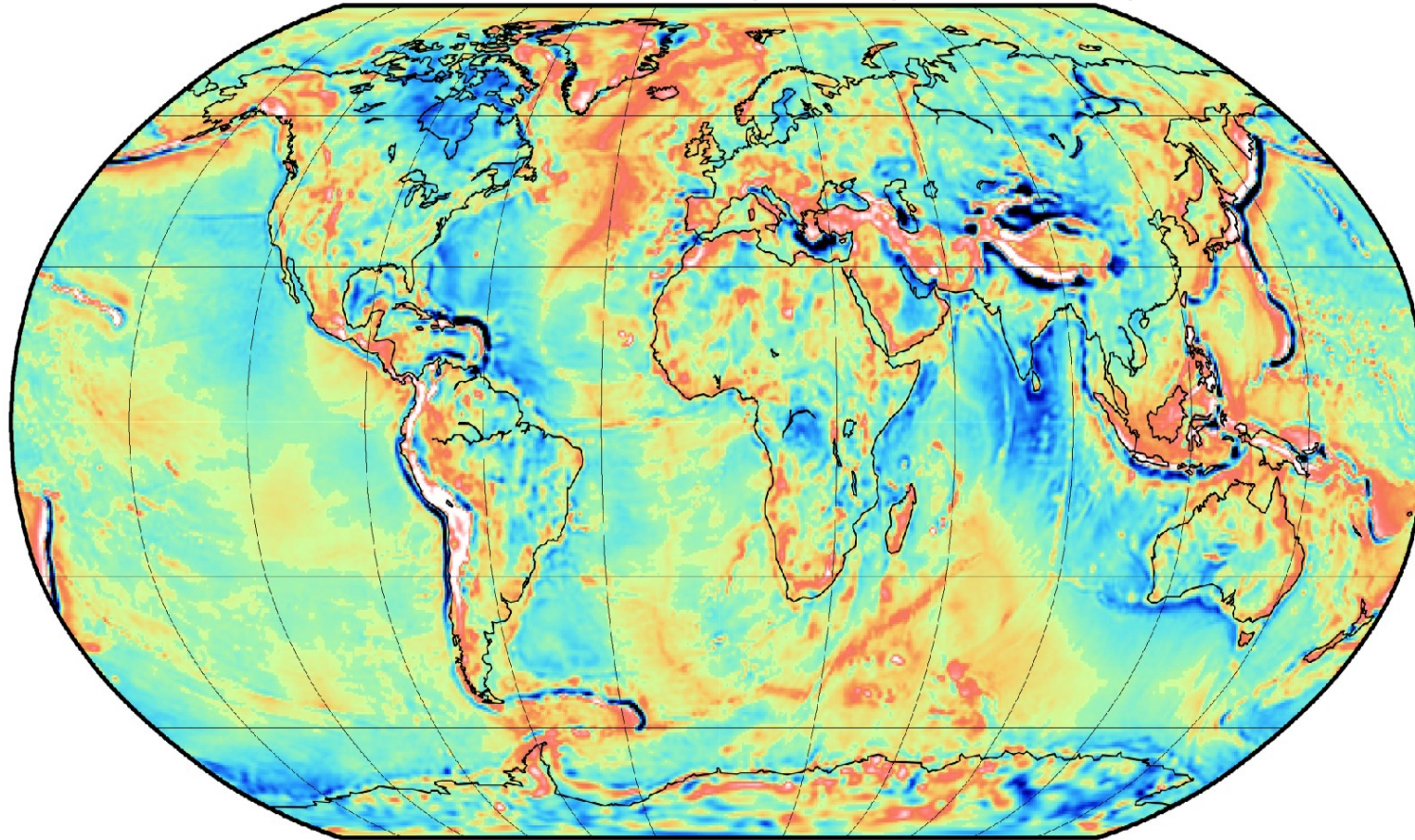


# *Back-up Slides*



# Combination Models

## EIGEN-GL05Cp gravity anomaly

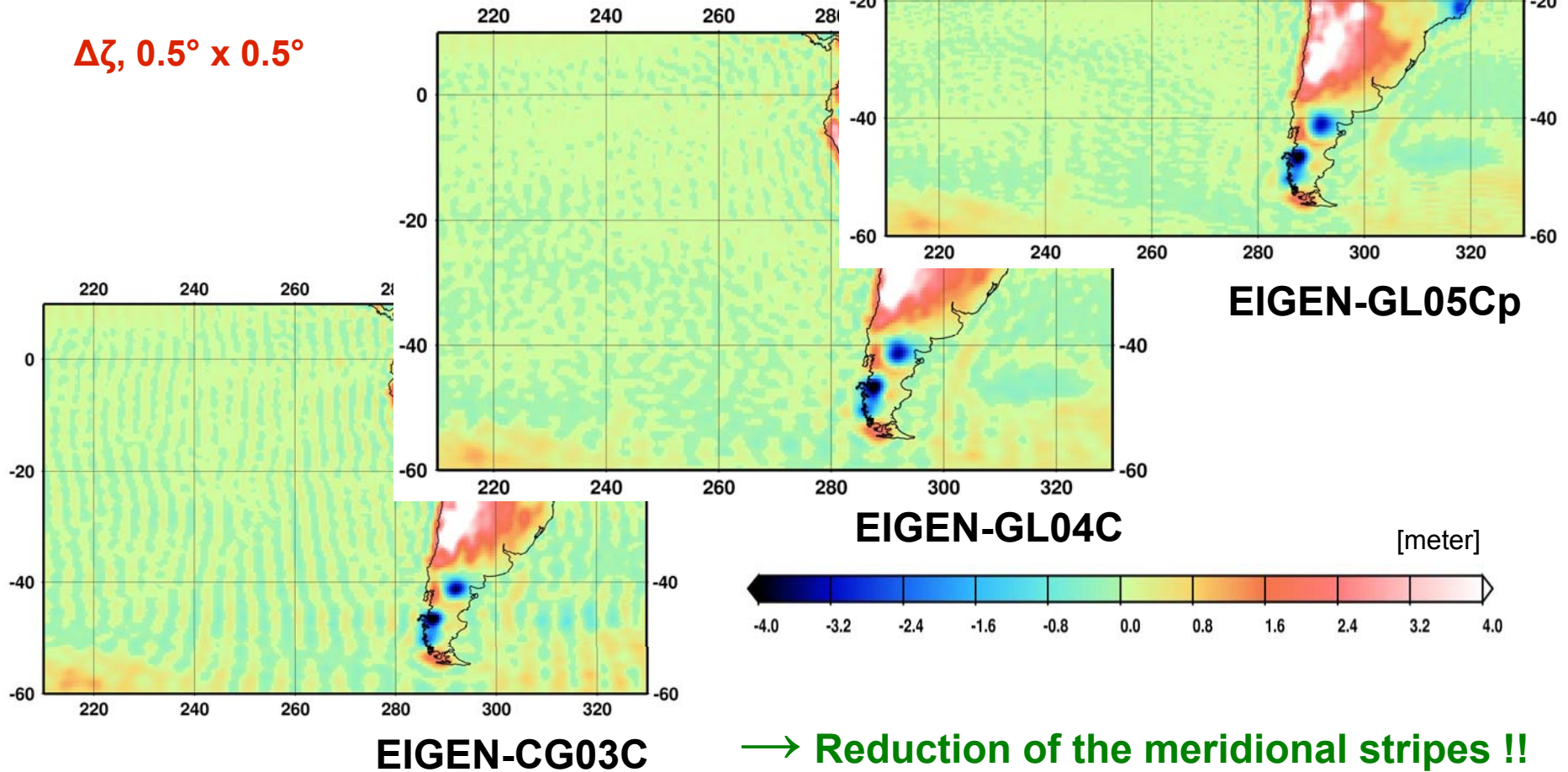


Increase of the spatial resolution by combination of satellite and surface data

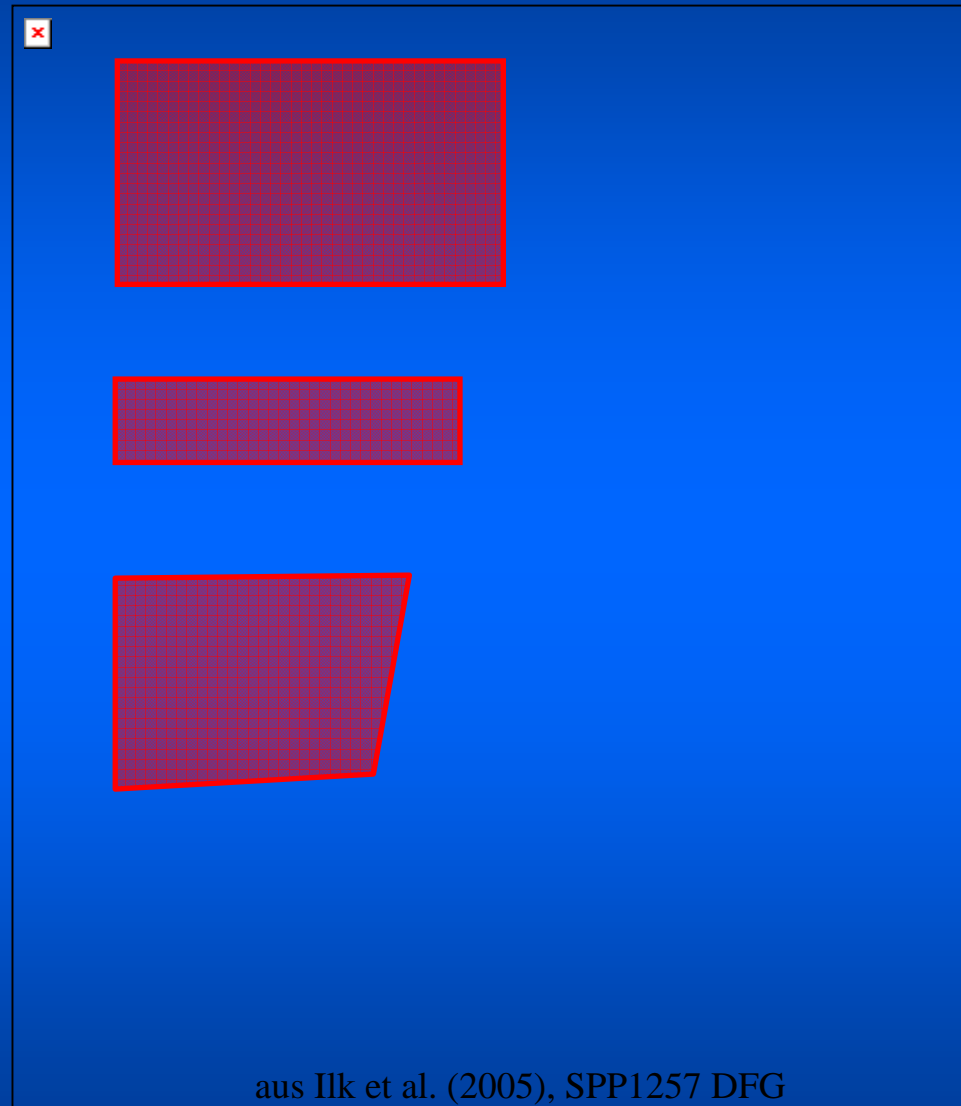
## Improvement of EIGEN-models

Geoid height differences vs. a global ground data only solution

$\Delta\zeta, 0.5^\circ \times 0.5^\circ$



# Conclusions



aus Ilk et al. (2005), SPP1257 DFG